



# How well are Southern Hemisphere teleconnection patterns predicted by seasonal climate models?

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**APEC Climate Symposium Lima/2008**

# Rossby Wave Theory

## Basic Theory – Rossby (1939, 1945)

The barotropic vorticity equation is:

$$\left( \frac{\partial}{\partial t} + U \frac{\partial}{\partial x} + V \frac{\partial}{\partial y} \right) \xi + V \beta_* = 0$$

$$U = \bar{U} + U'$$

Assuming that

$$V = V' \quad \xi = \xi'$$

And defining the perturbed streamfunction  $\psi$ , we have:

$$\left( \frac{\partial}{\partial t} + \bar{U} \frac{\partial}{\partial x} \right) \nabla^2 \psi + \beta_* \frac{\partial \psi}{\partial x} = 0$$

Assuming the wave solution

$$\psi = \text{Re} \left\{ A e^{i(kx + ly - \omega t)} \right\}$$

We get:

$$\omega = \bar{U}k - \frac{\beta_* k}{(k^2 + l^2)} \quad \text{or} \quad c_x = \bar{U} - \frac{\beta_*}{(k^2 + l^2)}$$

## Some characteristics of Rossby waves are:

- They propagate to the west
- They are dispersive

The group velocity is given by:

$$c_{g_x} = \frac{\partial \omega}{\partial k} = \frac{\omega}{k} + \frac{2\beta_* k^2}{(k^2 + l^2)^2} \quad \text{and} \quad c_{g_y} = \frac{\partial \omega}{\partial l} = \frac{2\beta_* k l}{(k^2 + l^2)^2}$$

For a stationary wave ( $\omega=0$ ;  $c=0$ ):

$$K^2 = (k^2 + l^2) = \frac{\beta_*}{U} = K_s^2$$

Playing with the equations, it is possible to define the ray path radius of curvature which is given by the simple expression

$$r = \frac{K_s^2}{\left[ k \frac{dK_s}{dy} \right]}$$

# Schematic $K_s$ profiles and ray path refraction

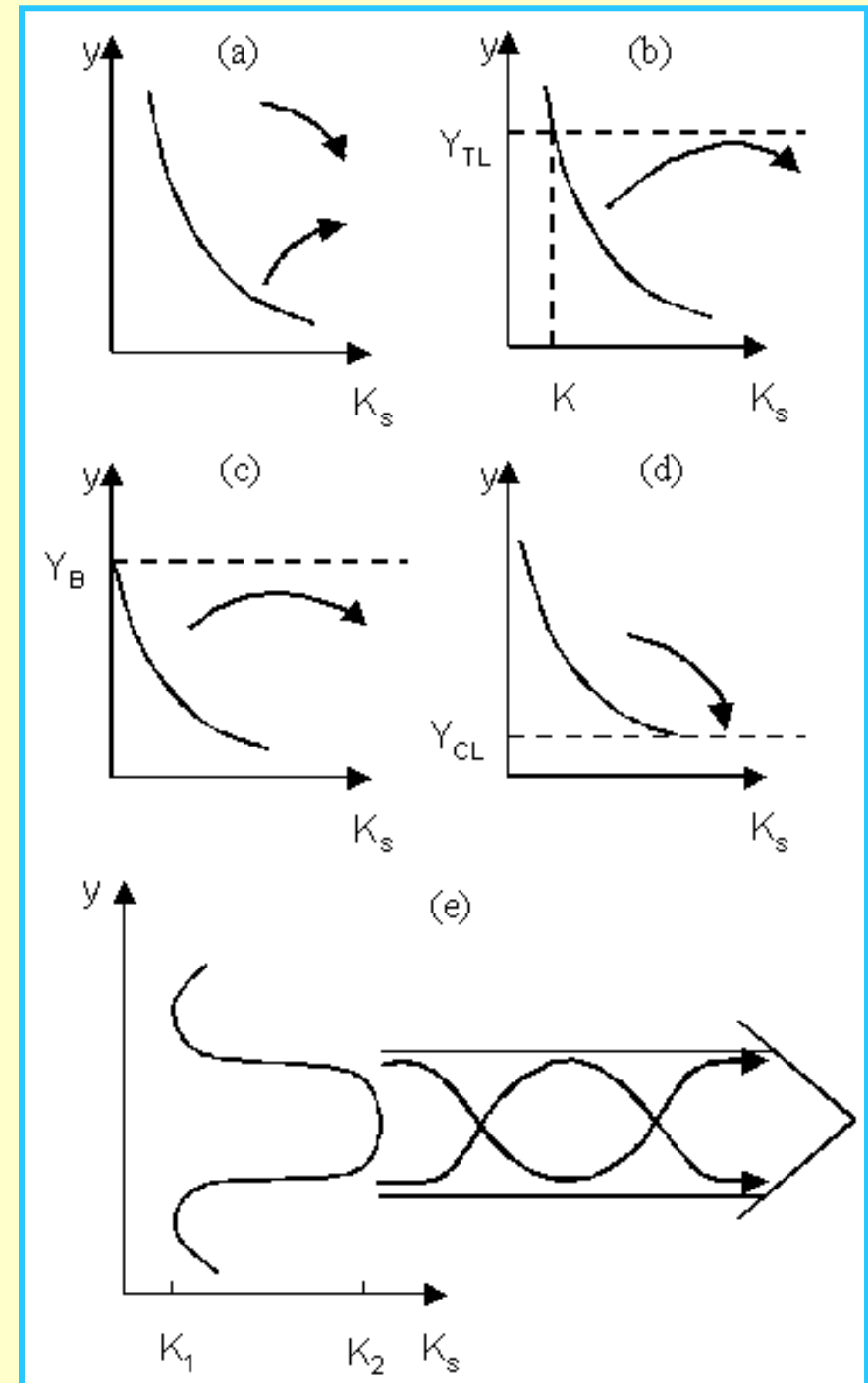
(a) Simple refraction

(b) Reflection from a turning latitude  $Y_{TL}$  at which  $K_s = k$

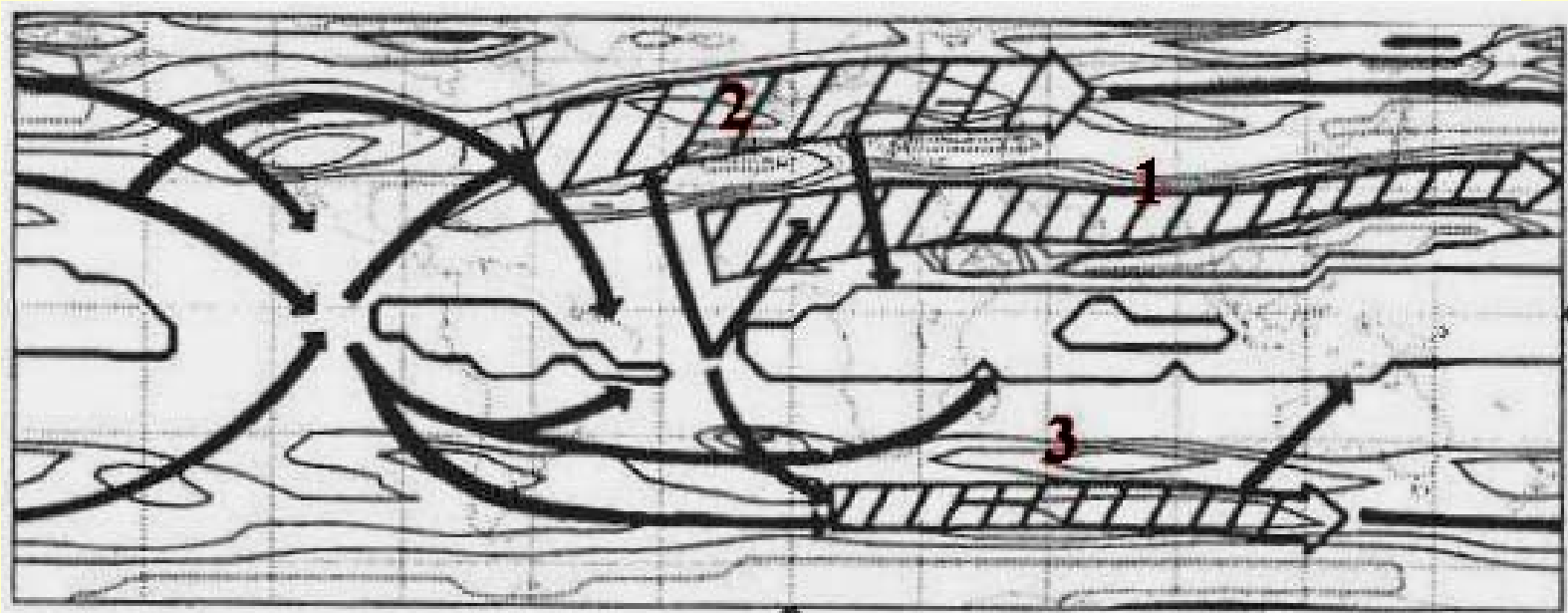
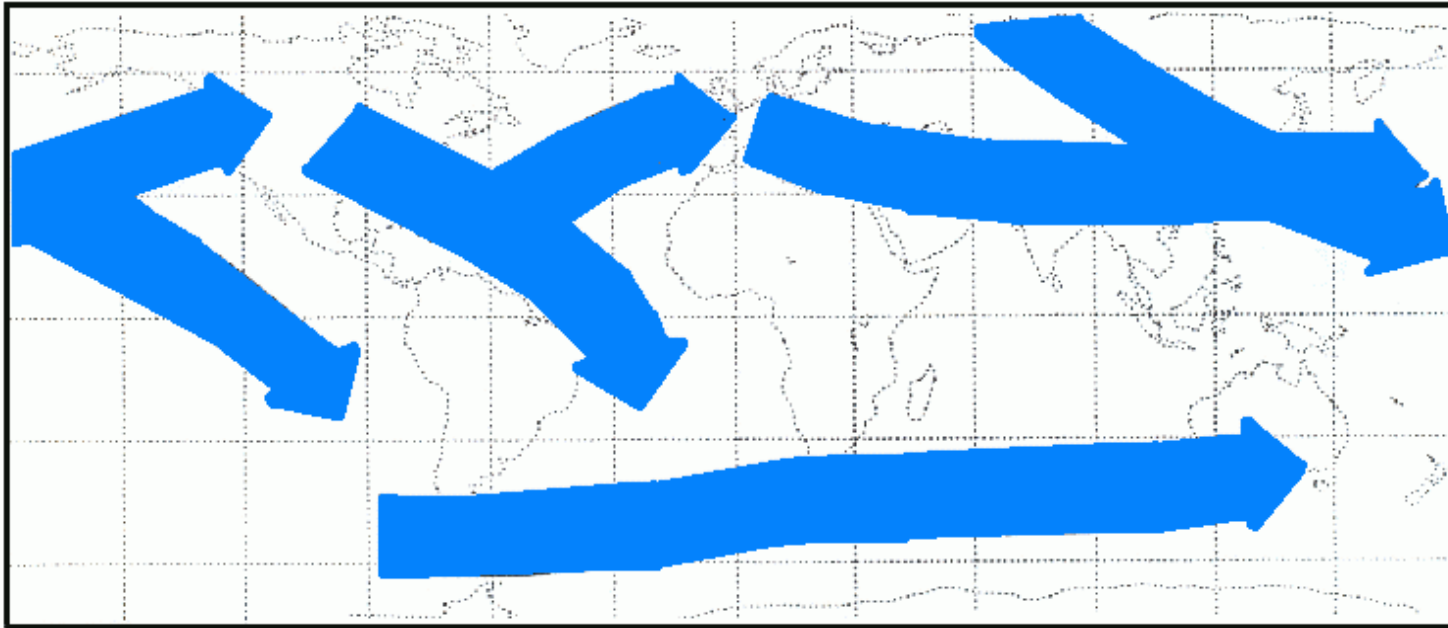
(c) Reflection of all wavenumbers before a latitude  $Y_B$  at which  $\beta^* = 0$

(d) Refraction into a critical latitude  $Y_{CL}$  at which  $U = 0$

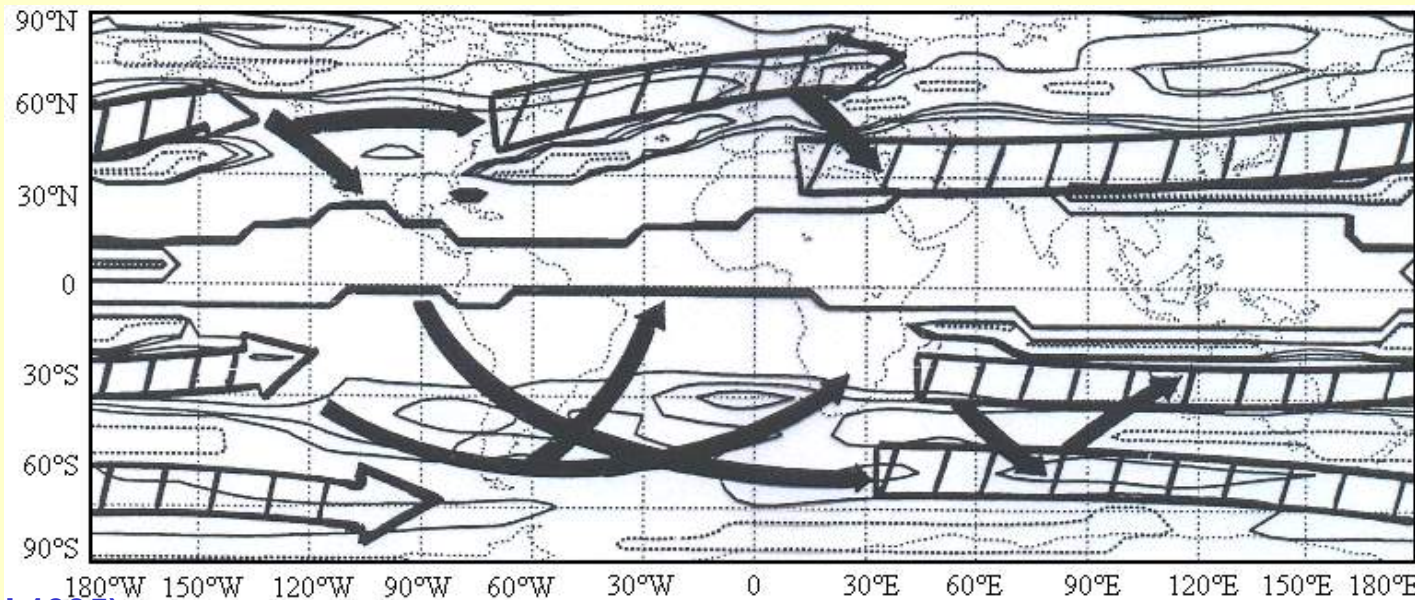
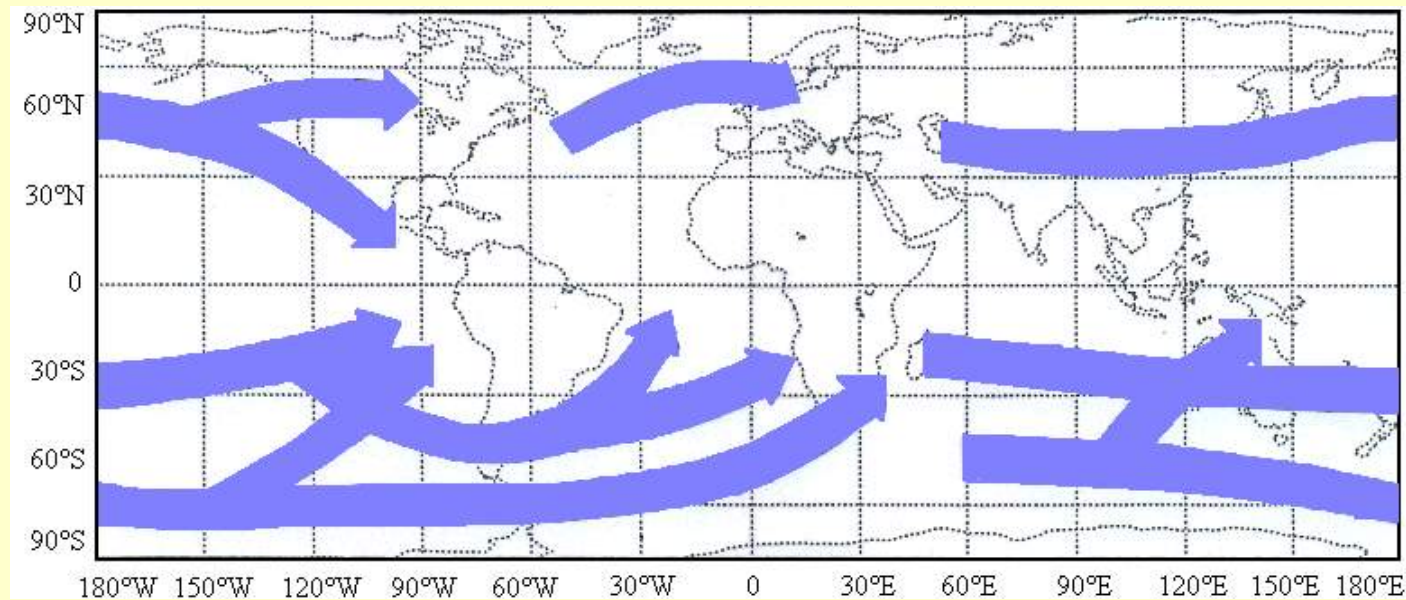
(e) waveguide effect of a  $K_s$  maximum.



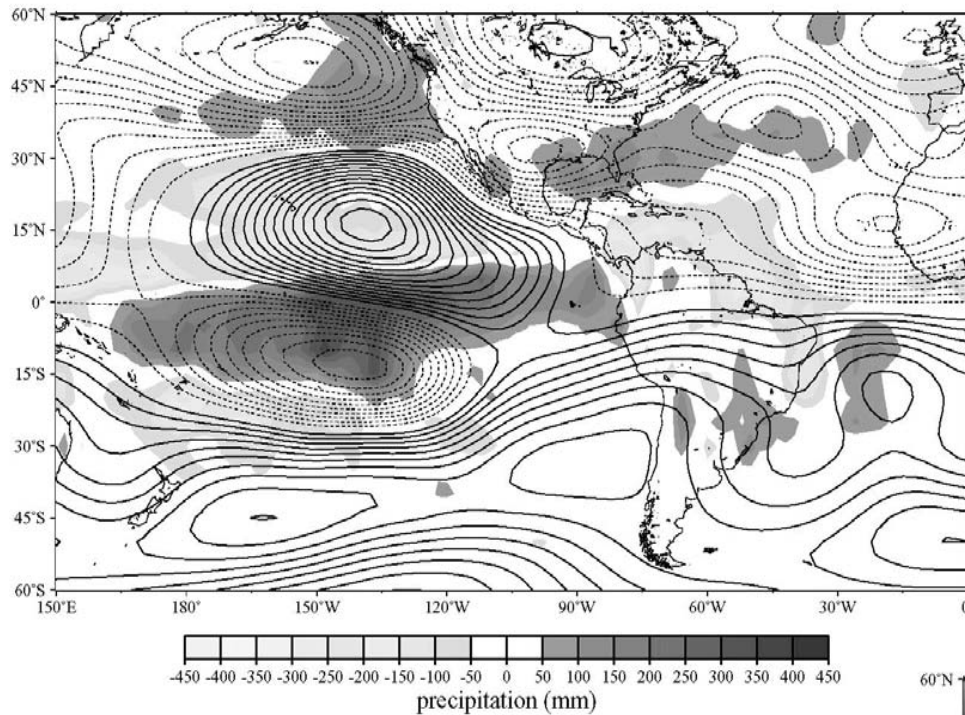
# Main teleconnection patterns obtained from observational analysis and numerical modeling - DJF



# Main teleconnection patterns obtained from observational analysis and numerical modeling - JJA



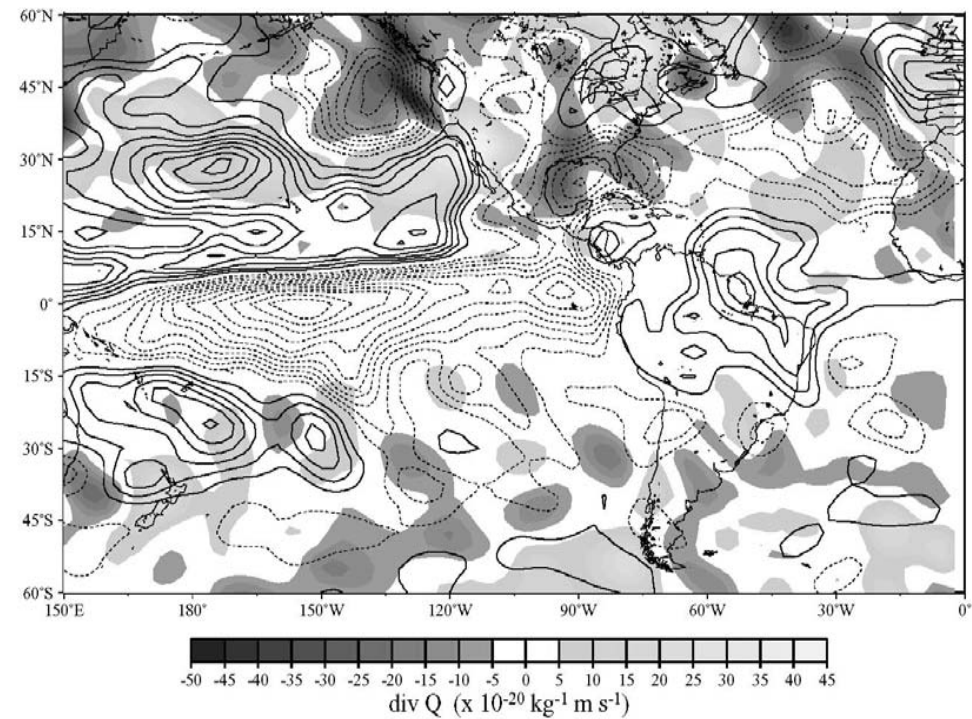
(Ambrizzi et al 1995)



# Stream Function, Omega, Precip and Q vector - average for four El Niño events

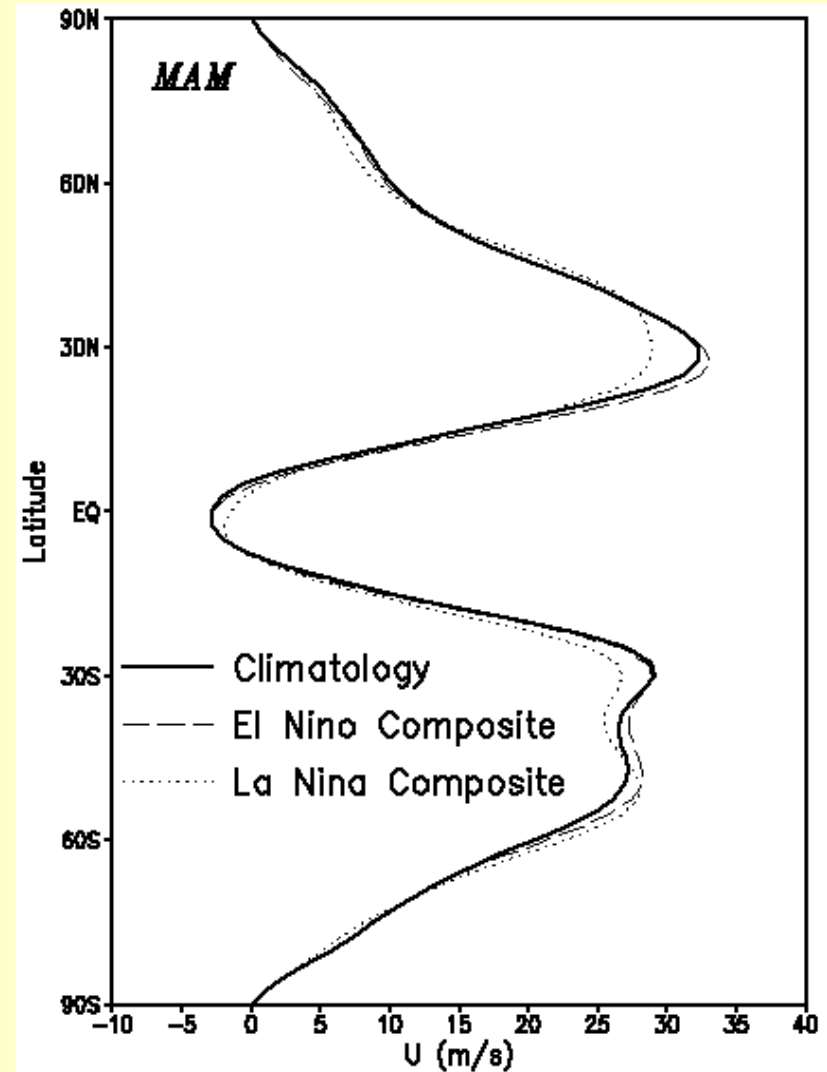
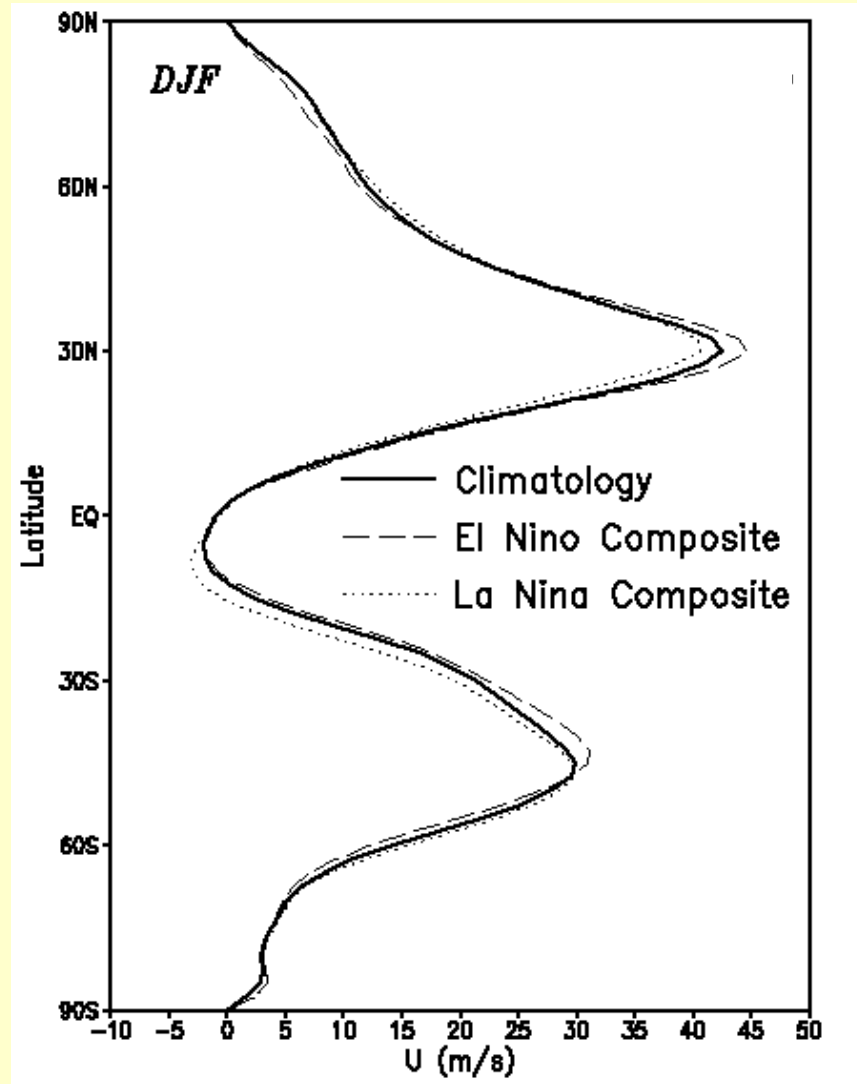
Div of Q anomaly of omega

Anomaly of  $\Psi$  and precipitation



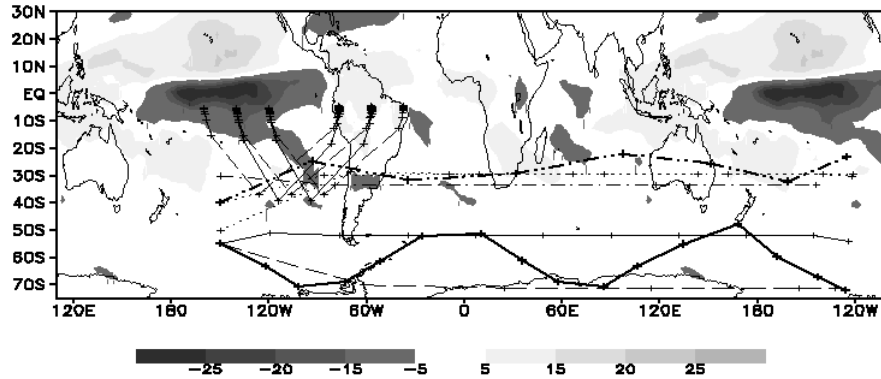
(Magaña e Ambrizzi, 2005)

# ZONAL WIND - U

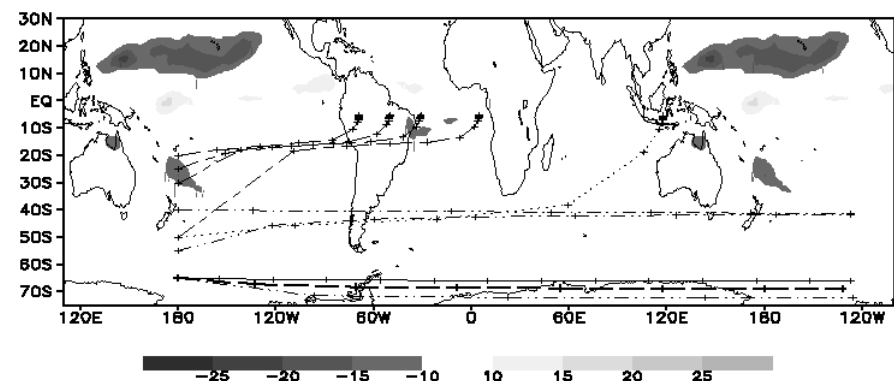
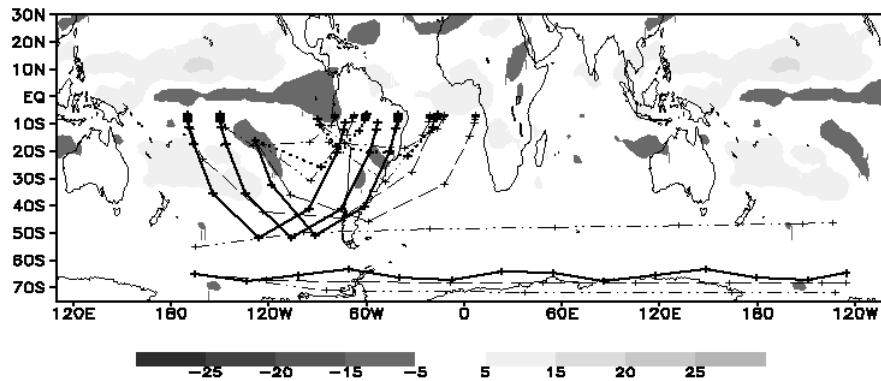
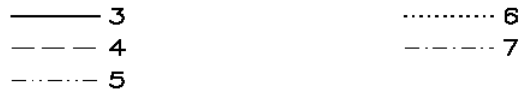
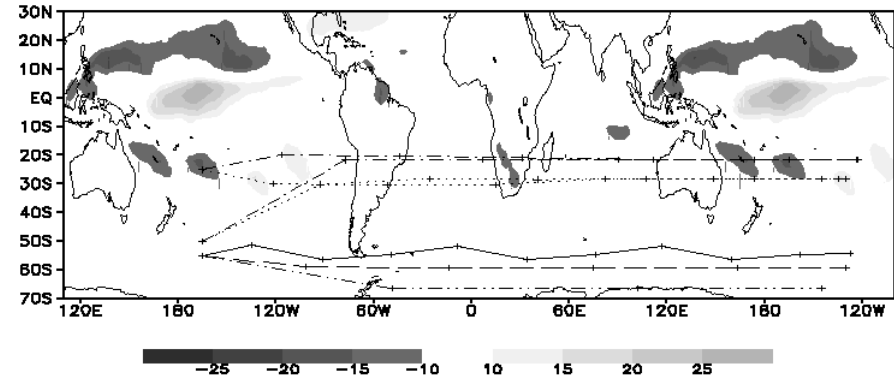


# RAY TRACING FOR EL NIÑO/LA NIÑA

(EN/DJF)



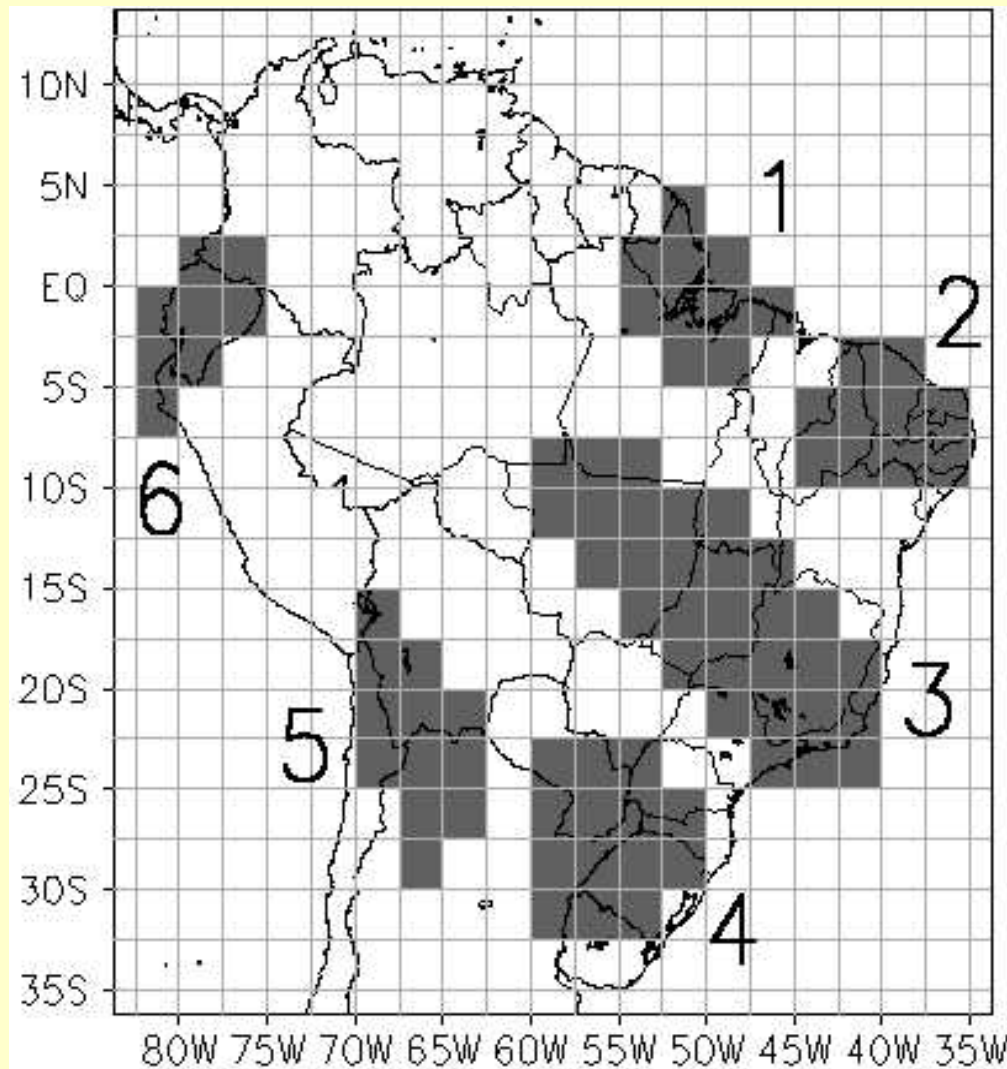
(LN/DJF)



(EN/MAM)

(LN/MAM)

# ENSO episodes and the South American Regional Precipitation



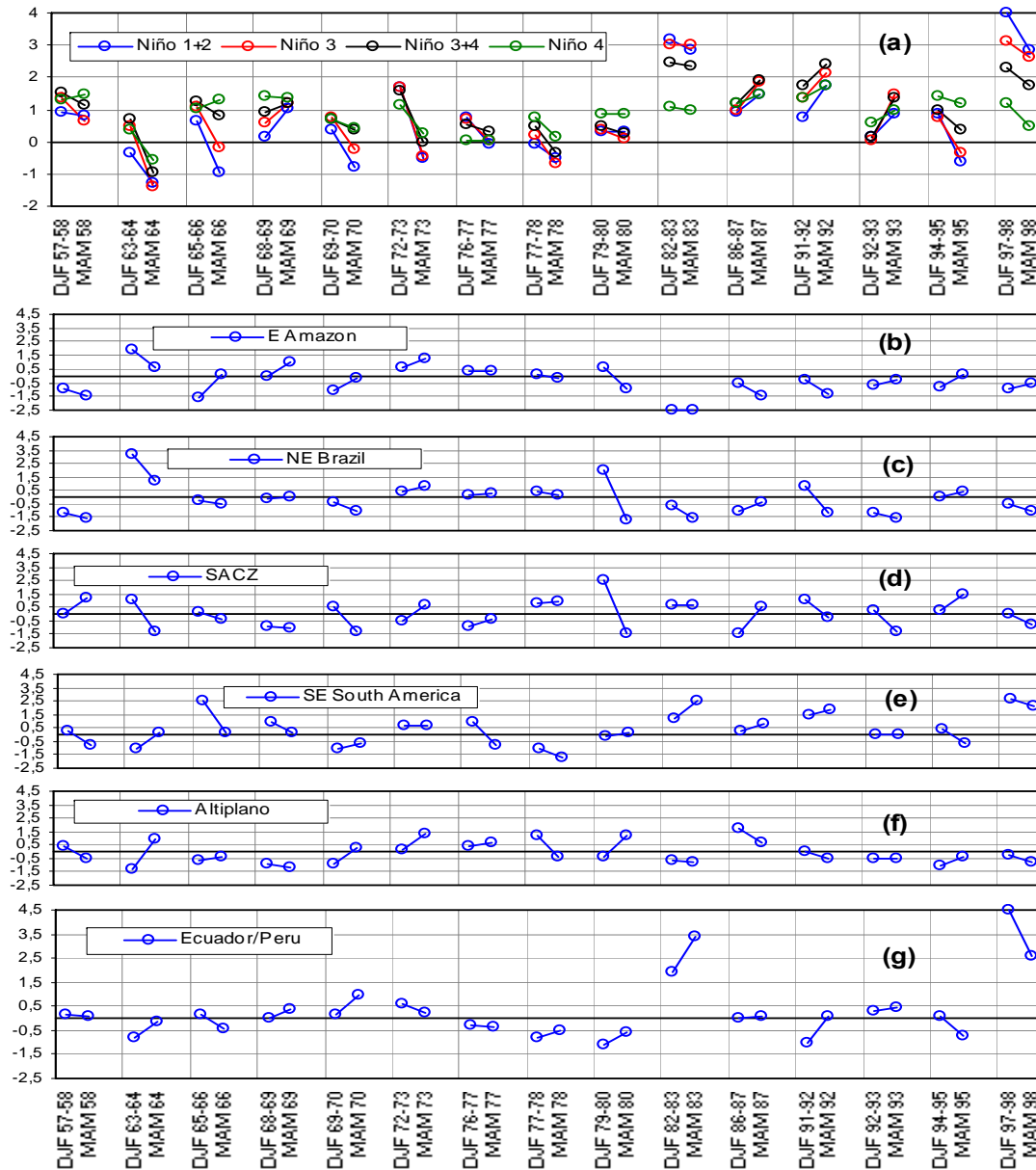
Seasonal Standardized index for the austral summer (DJF) and autumn (MAM)

$$I = (X_i - \bar{X}_i) / \sigma_i$$

were calculated for the precipitation over key-areas of South America and the SSTa over the Niño 1.2, 3, 3.4 and 4

(Ambrizzi and Souza, 2003)

# El Niño Episodes and regional precipitation over South America during the period 1950-1990



aSST Niños

E Amazônia

NE Brasil

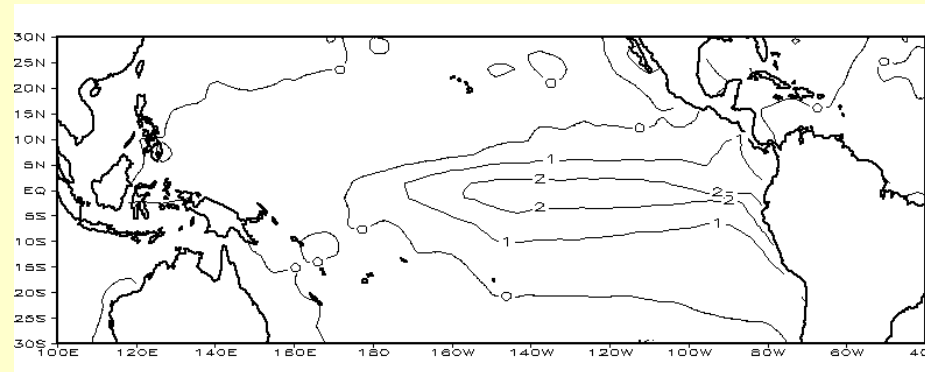
SACZ

SE South America

Altiplano

Equator Peru

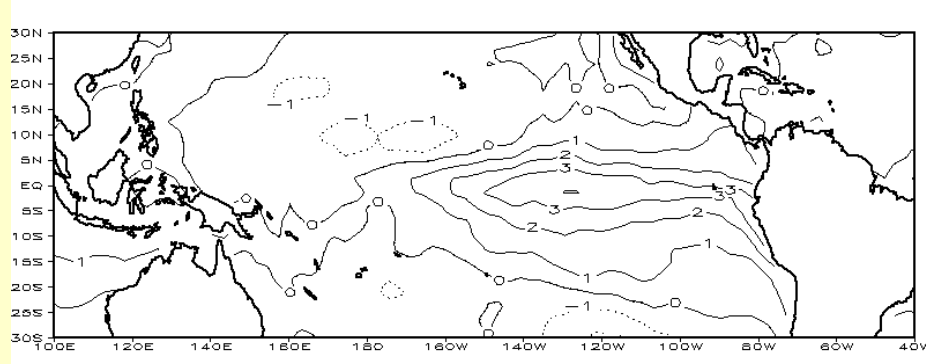
# SSTs for 4 EL NIÑO EVENTS



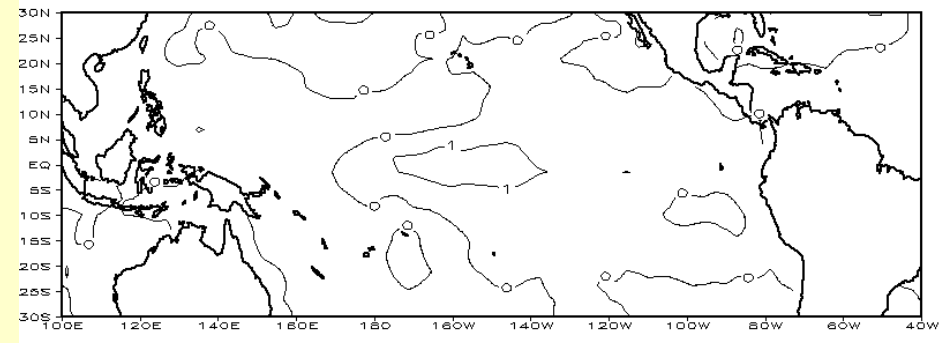
Average for 4 events

El Niño 82/83

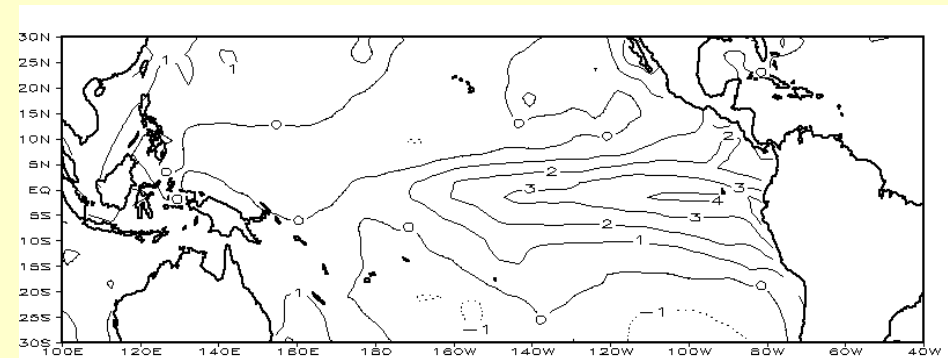
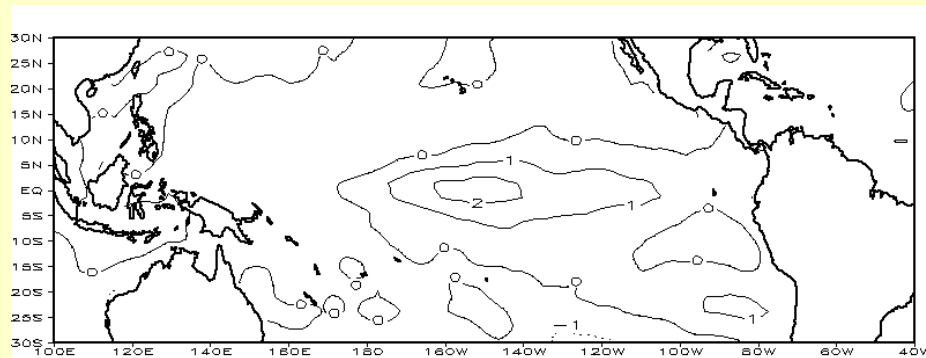
El Niño 86/87



El Niño 91/92

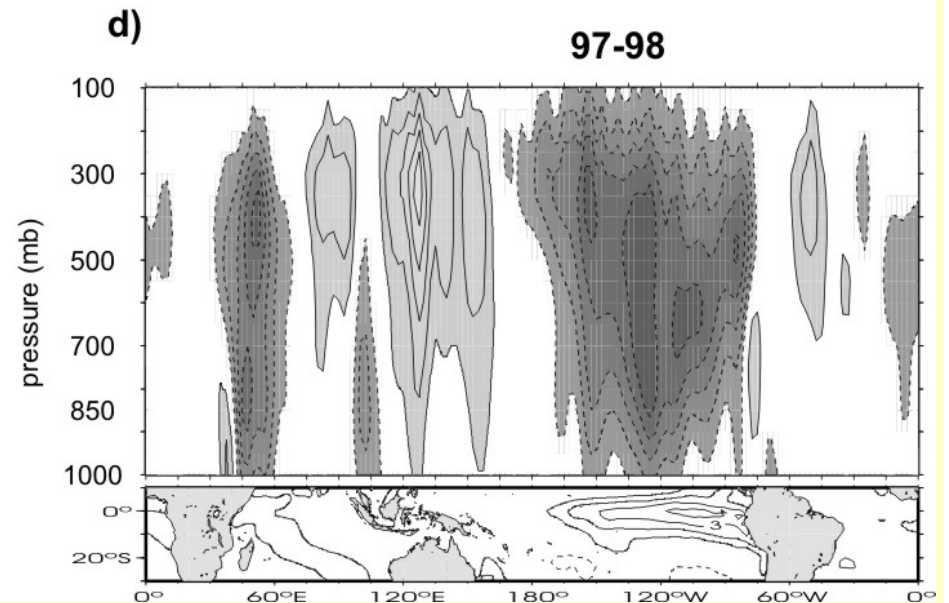
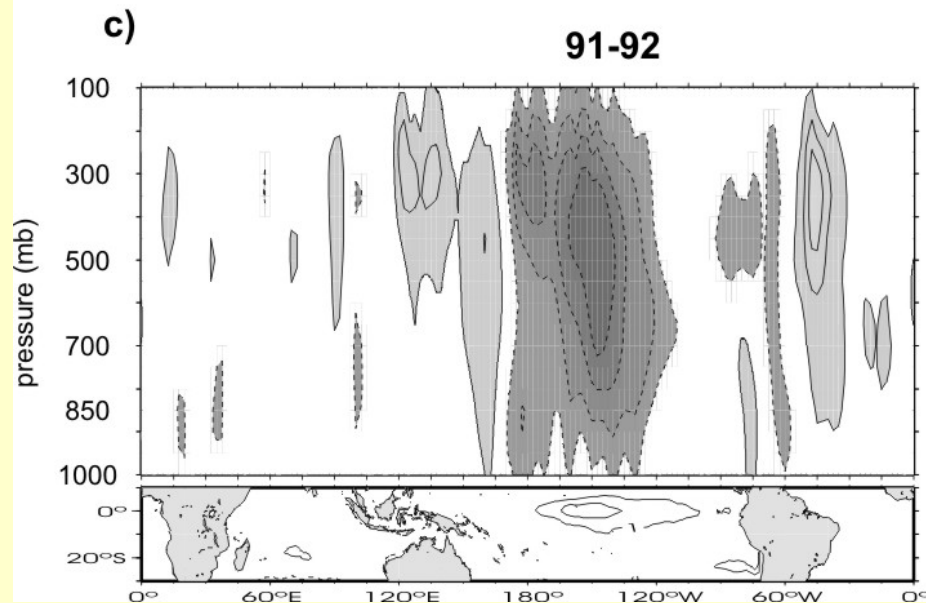
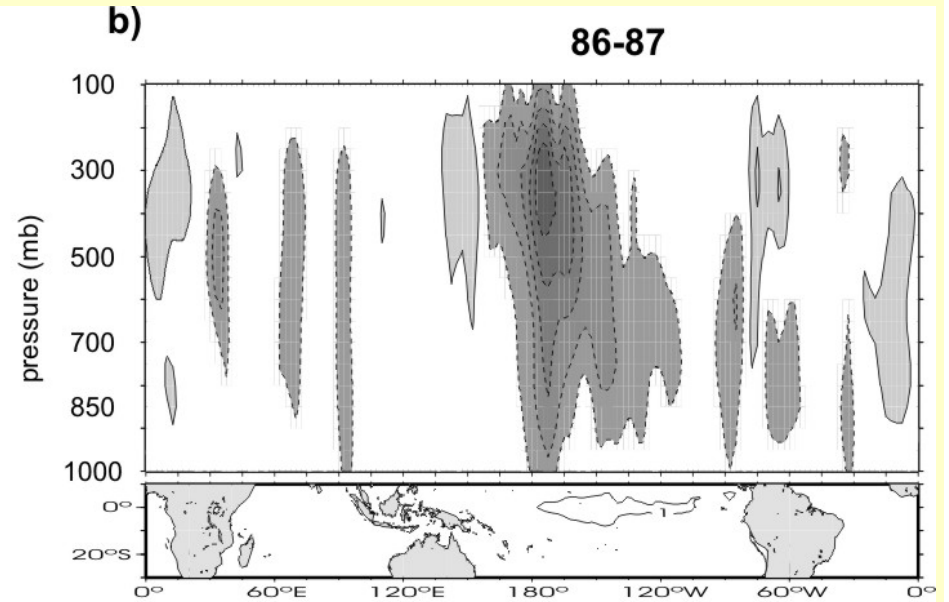
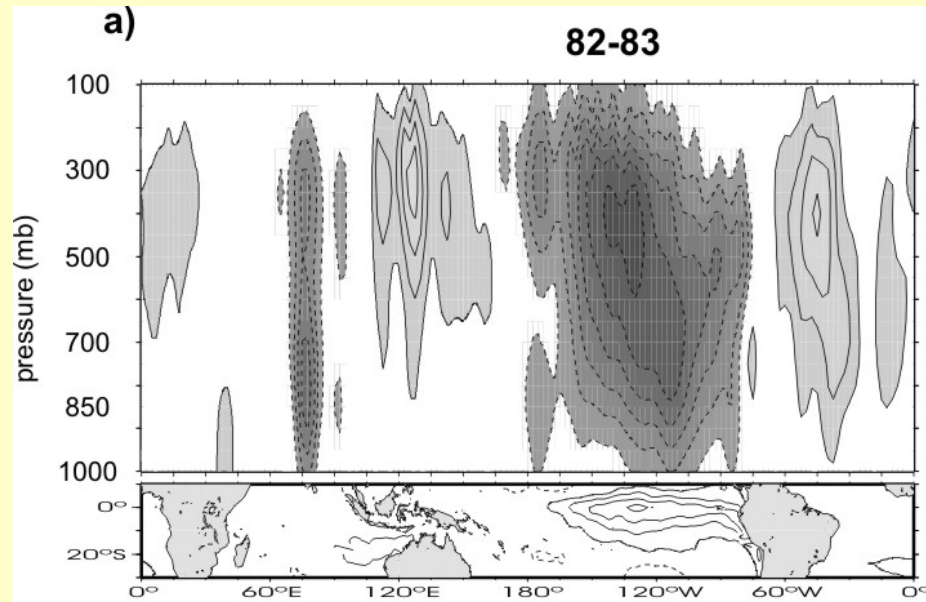


El Niño 97/98



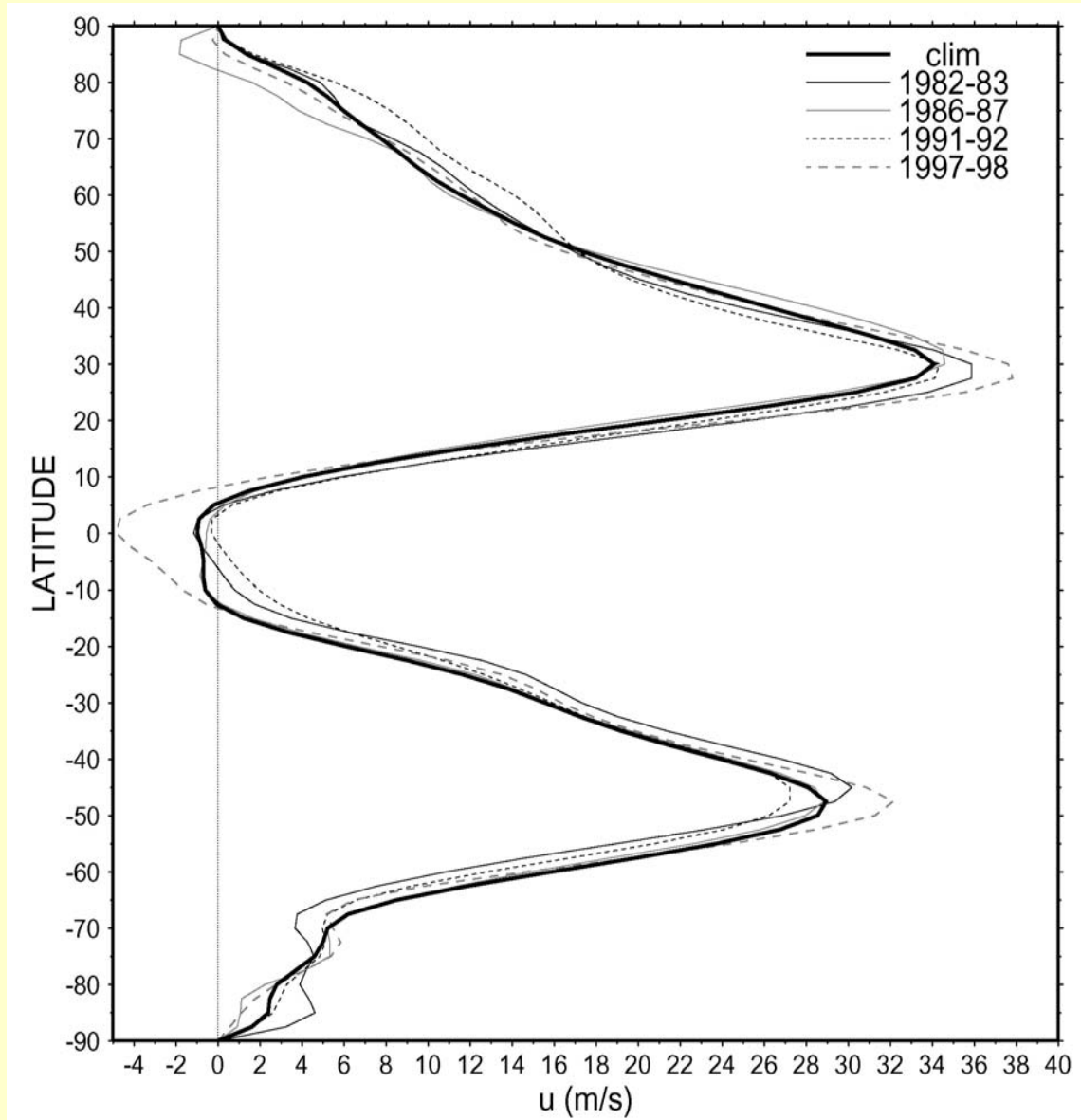
(Drumond e Ambrizzi, 2003)

# OMEGA ANOMALY

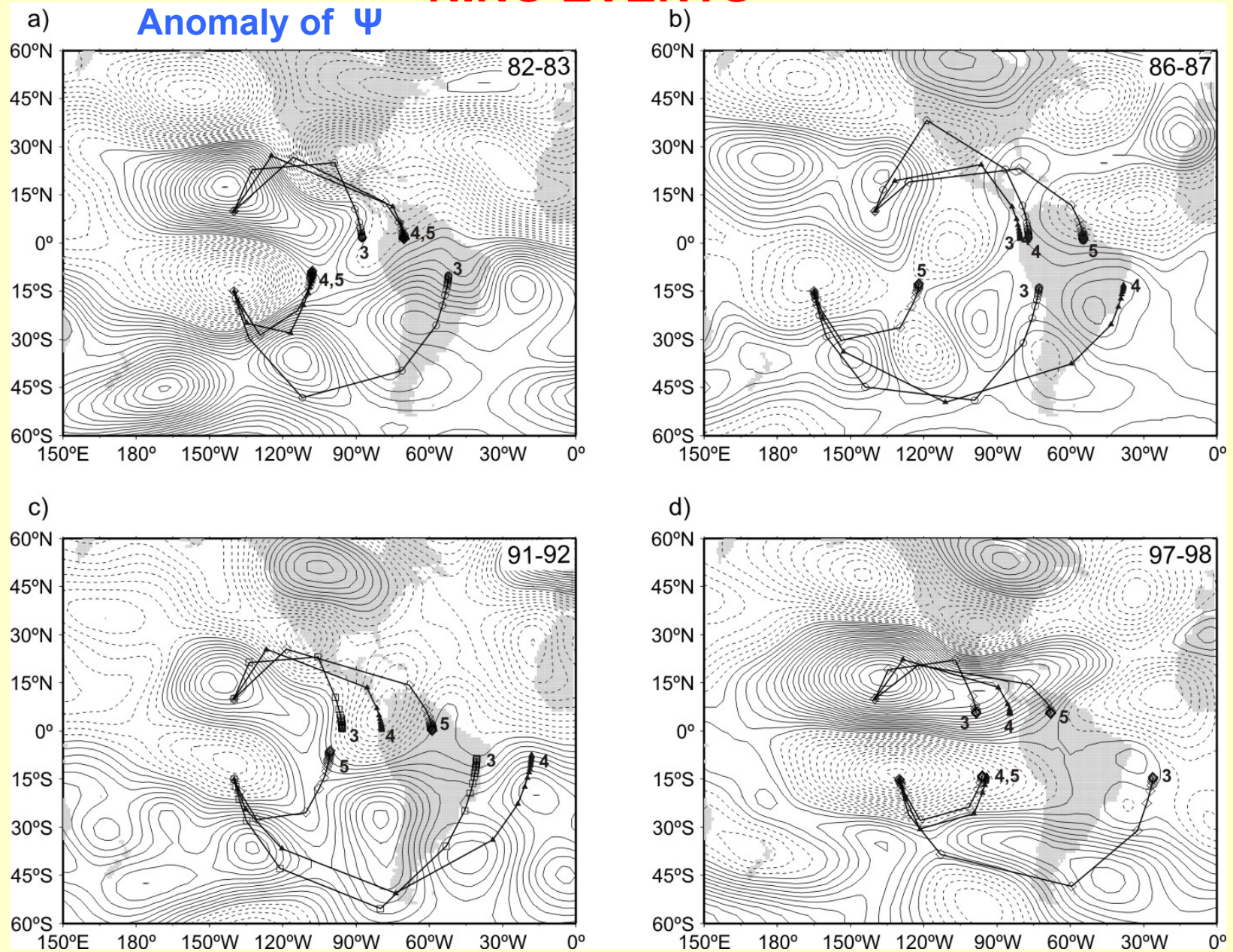


(Magaña e Ambrizzi 2005)

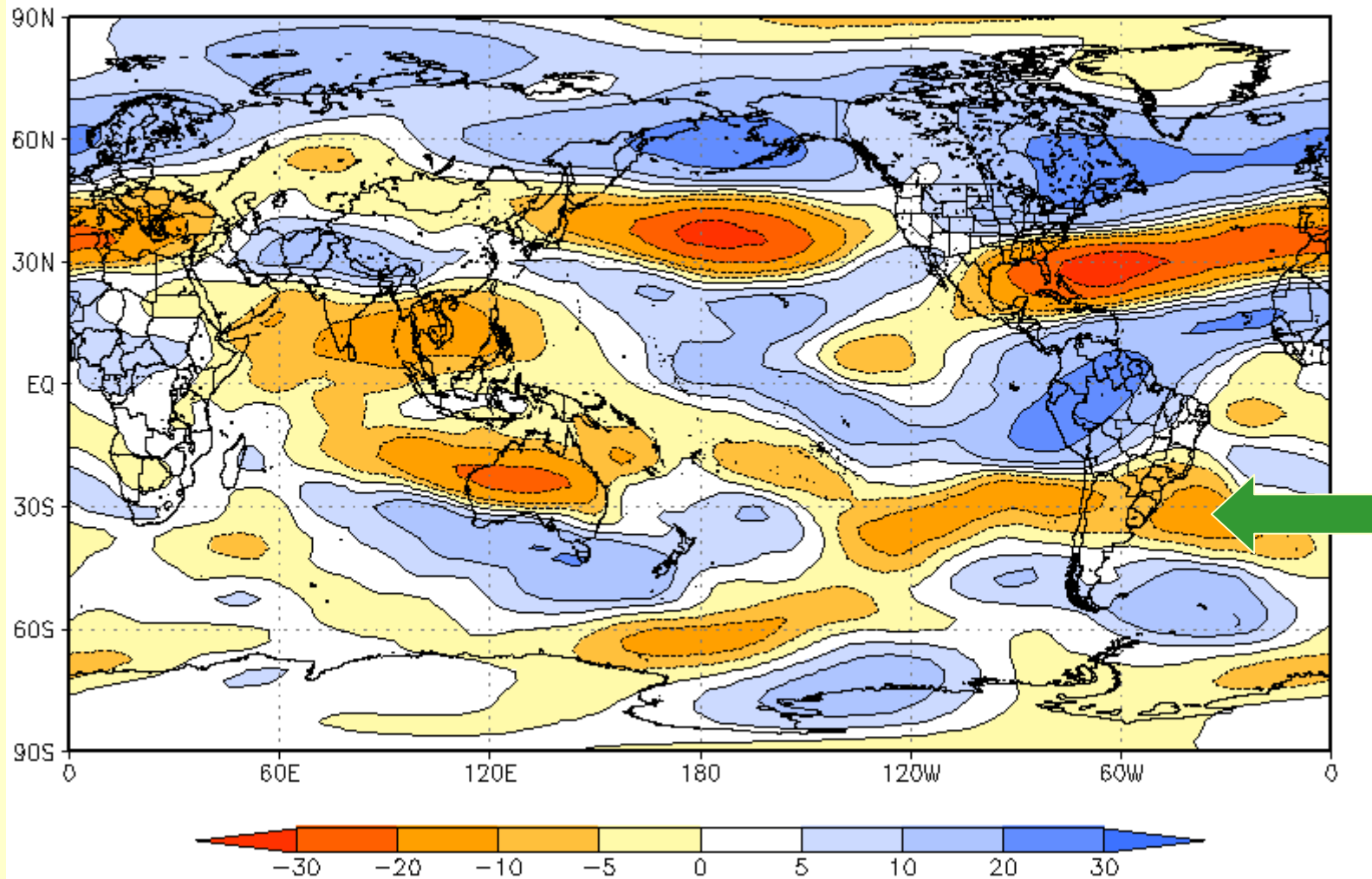
# MEAN ZONAL WIND, 250 hPa, DJF



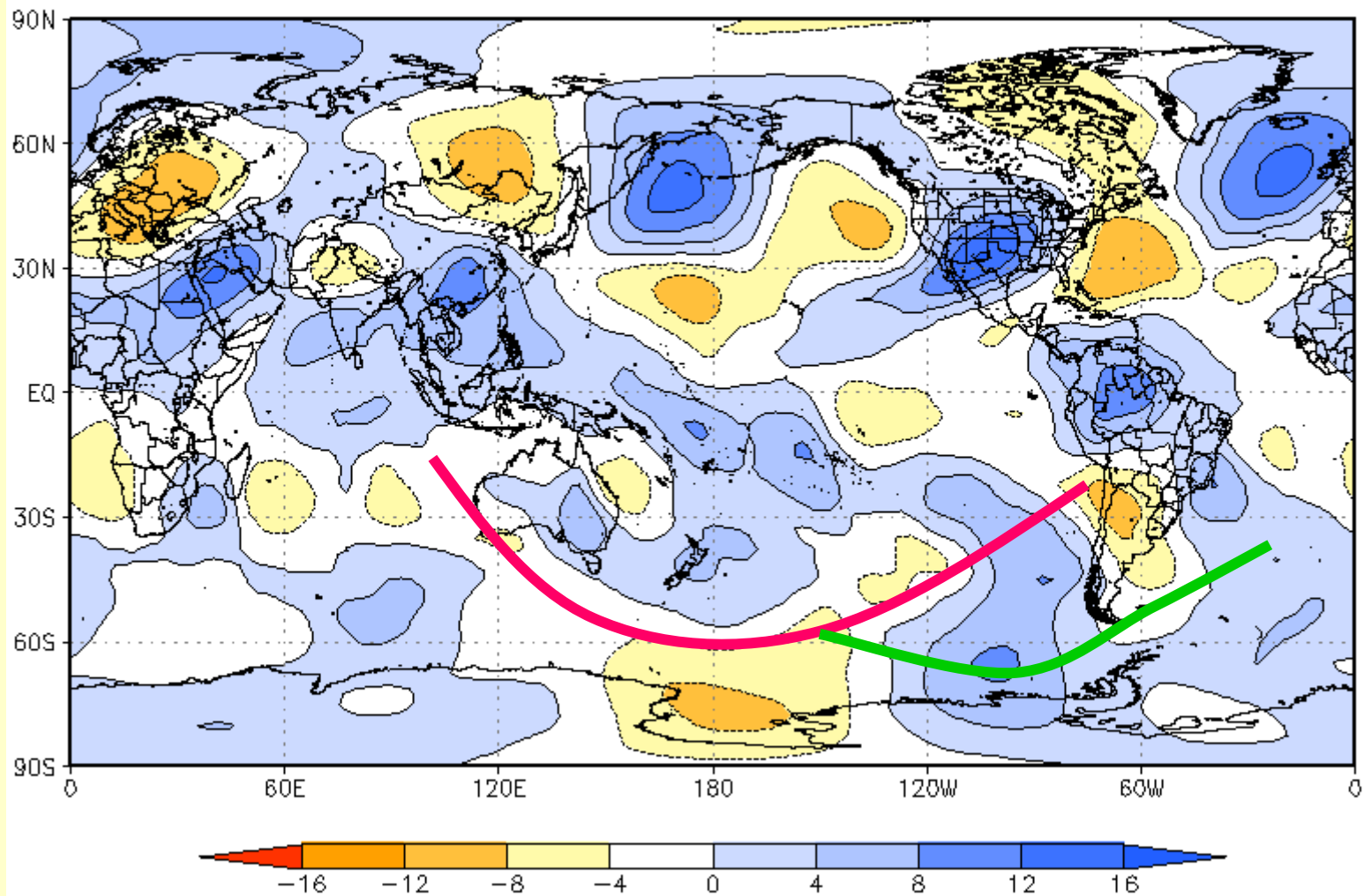
# STREAMFUNCTION AND RAY TRACING FOR 4 EL NIÑO EVENTS



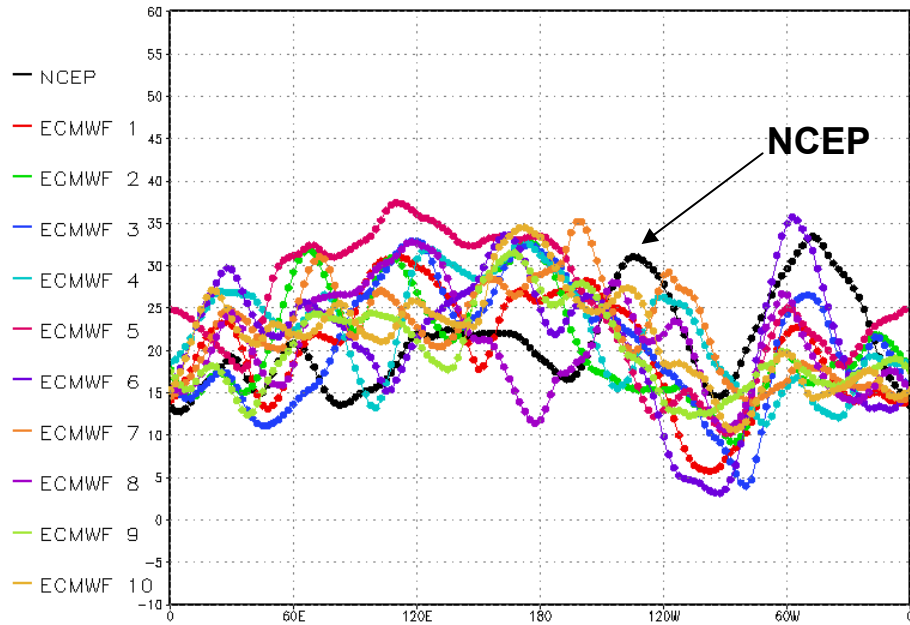
Diference Zonal Wind (200 mb)  
ECMWF 1 – NCEP (March/1981)



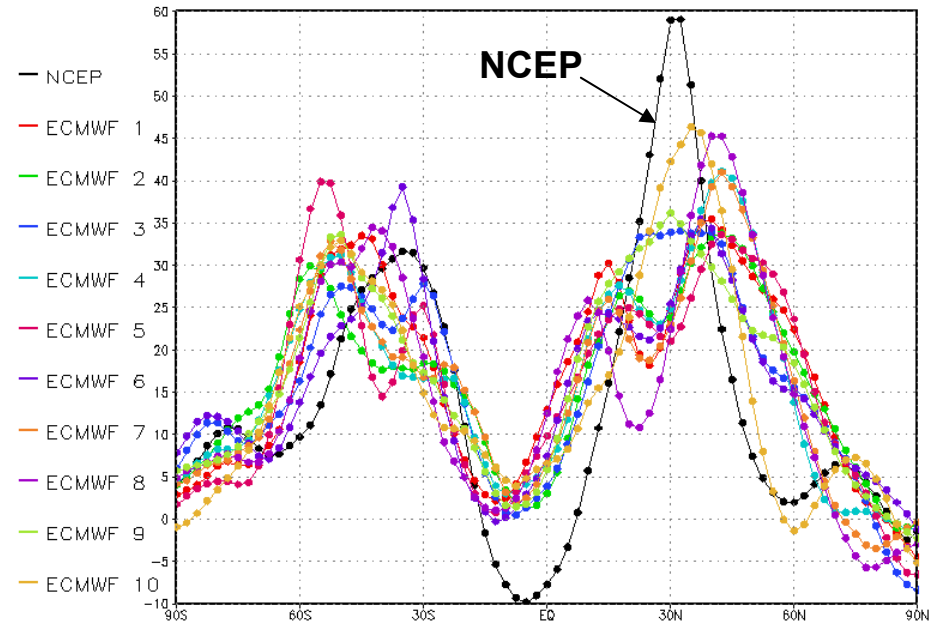
Meridional Wind – ECMWF 1 – 200 mb – March/1981



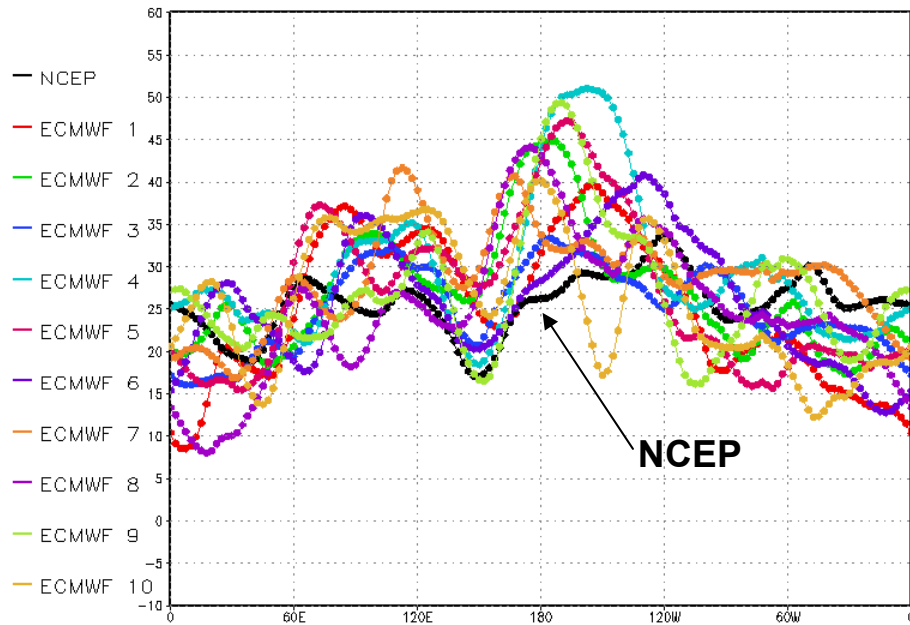
Cross Section 32°S – Zonal Wind 200mb (March/81)



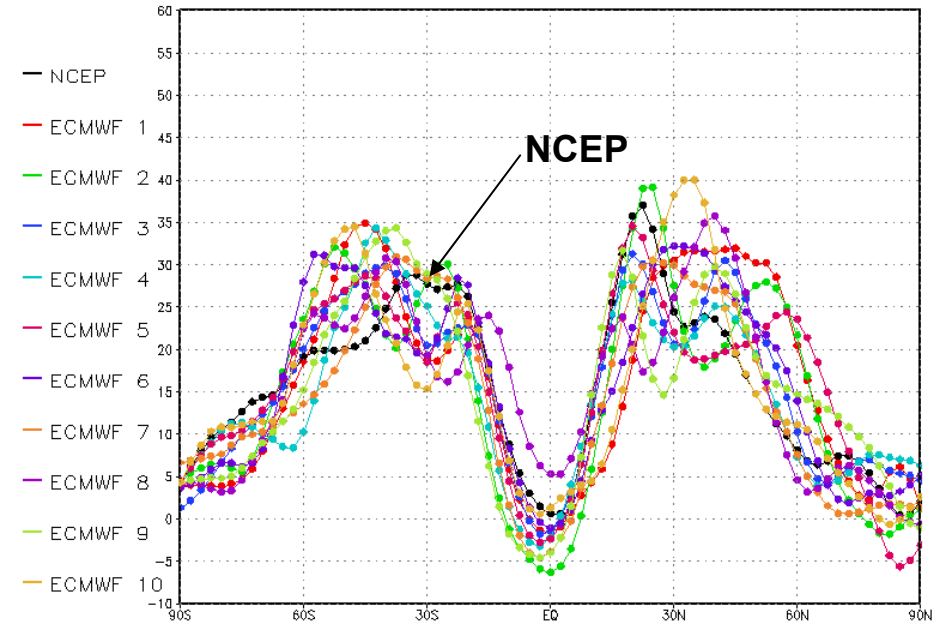
Cross Section 54°W – Zonal Wind 200mb (March/81)



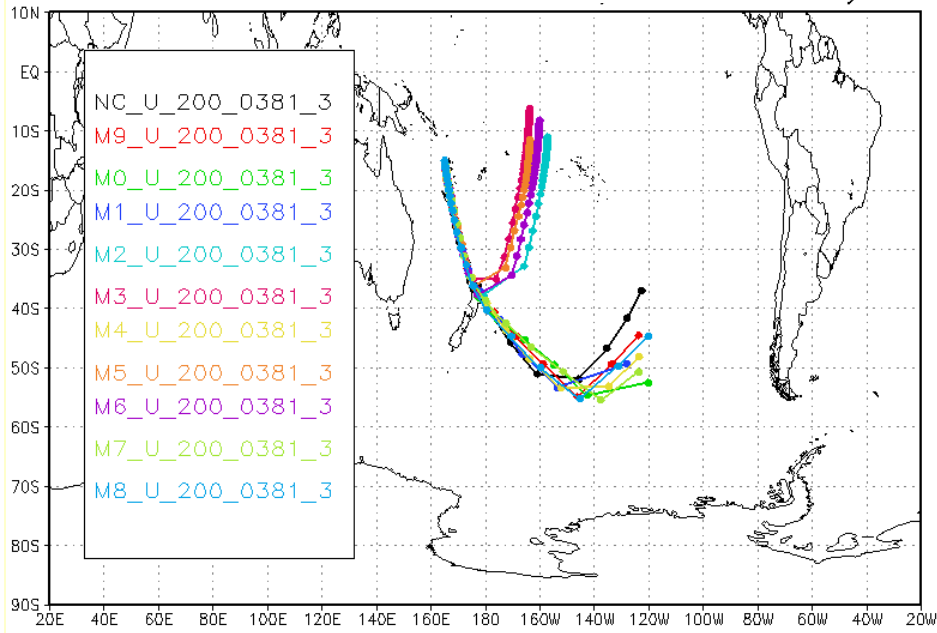
Cross Section 32°S – Zonal Wind 200mb (April/81)



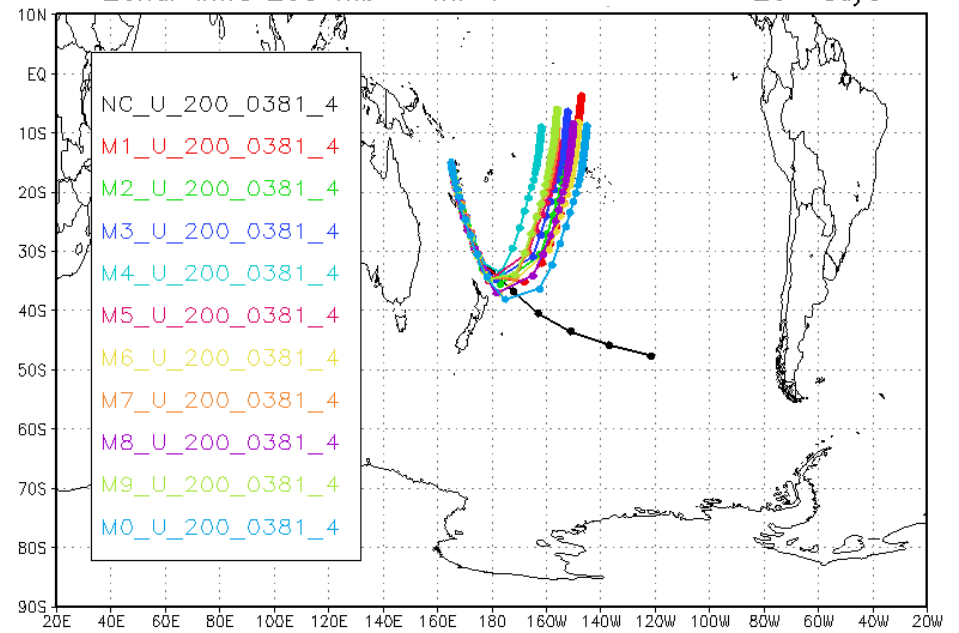
Cross Section 54°W – Zonal Wind 200mb (April/81)



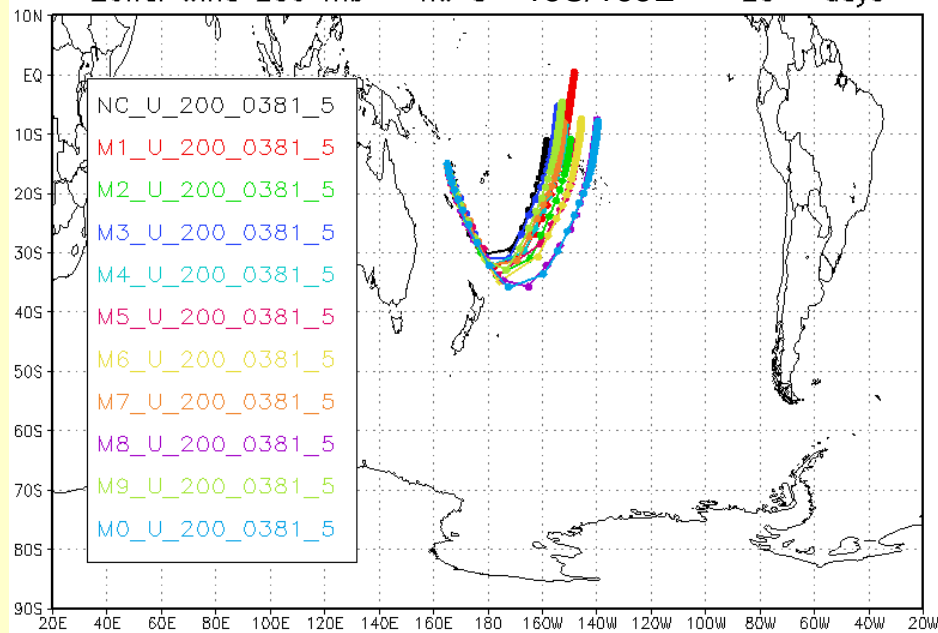
Zonal wind 200 mb - nw 3 - 15S/165E - 20 days



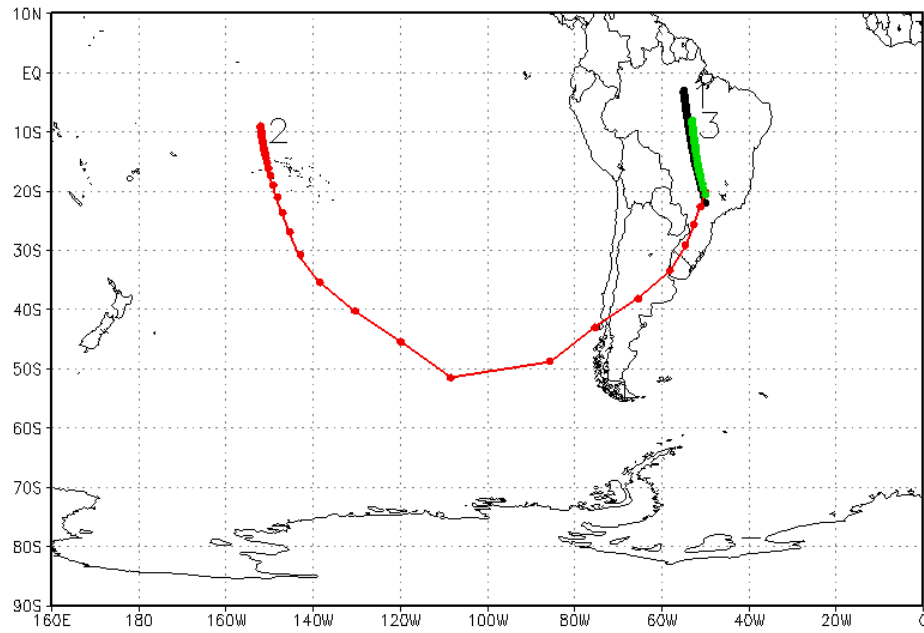
Zonal wind 200 mb - nw 4 - 15S/165E - 20 days



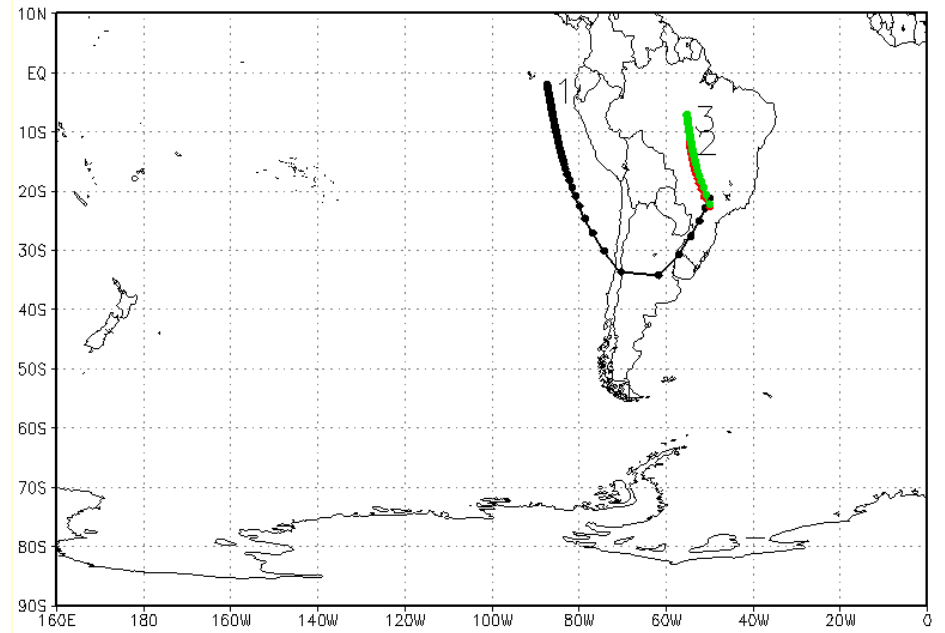
Zonal wind 200 mb - nw 5 - 15S/165E - 20 days



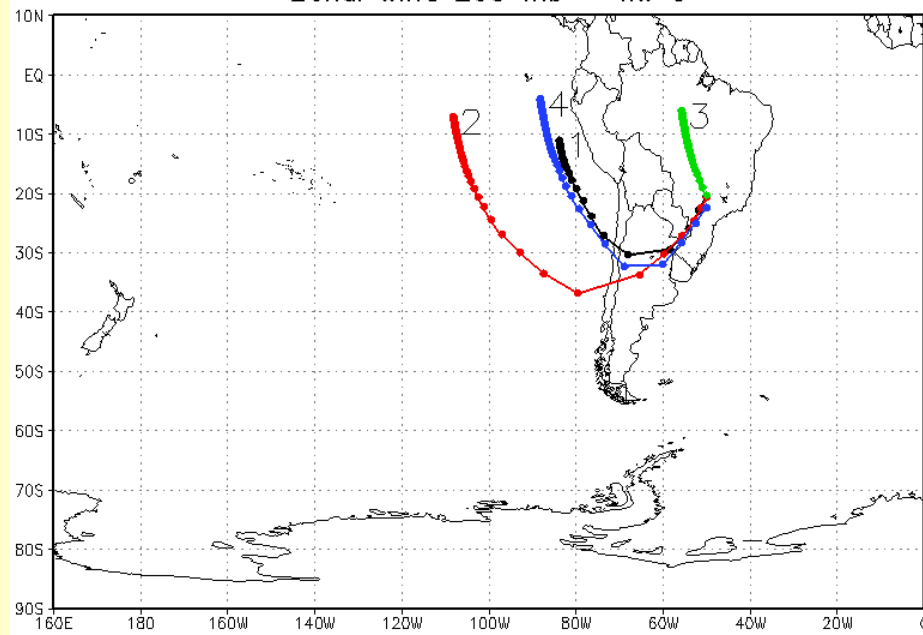
Zonal wind 200 mb - nw 3



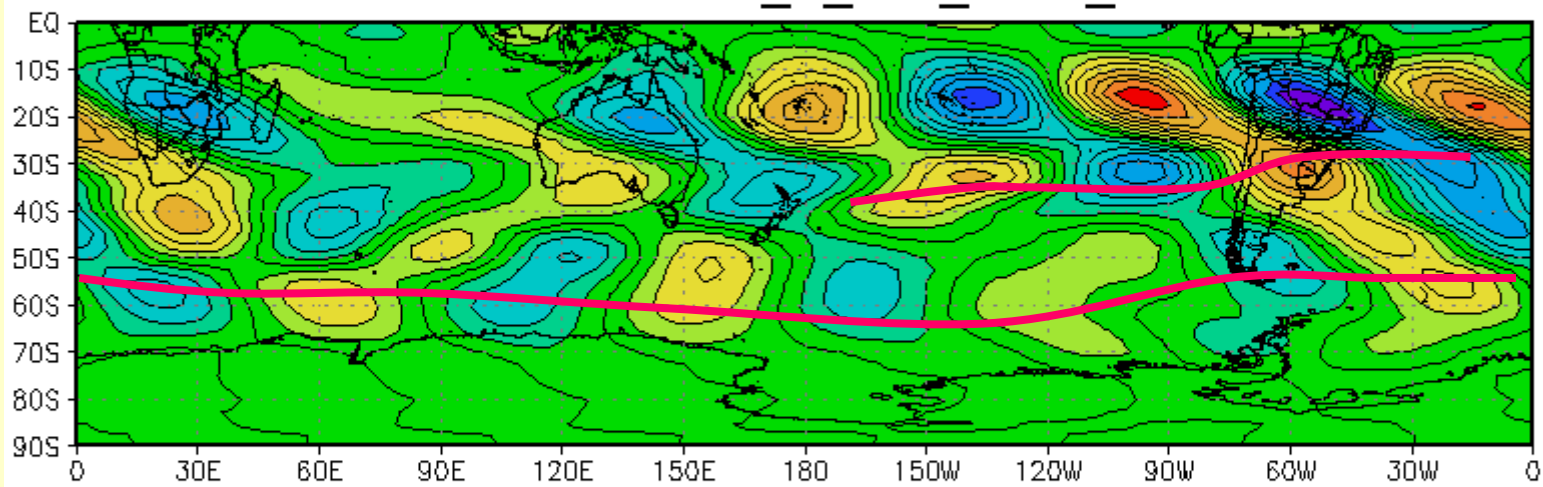
Zonal wind 200 mb - nw 4



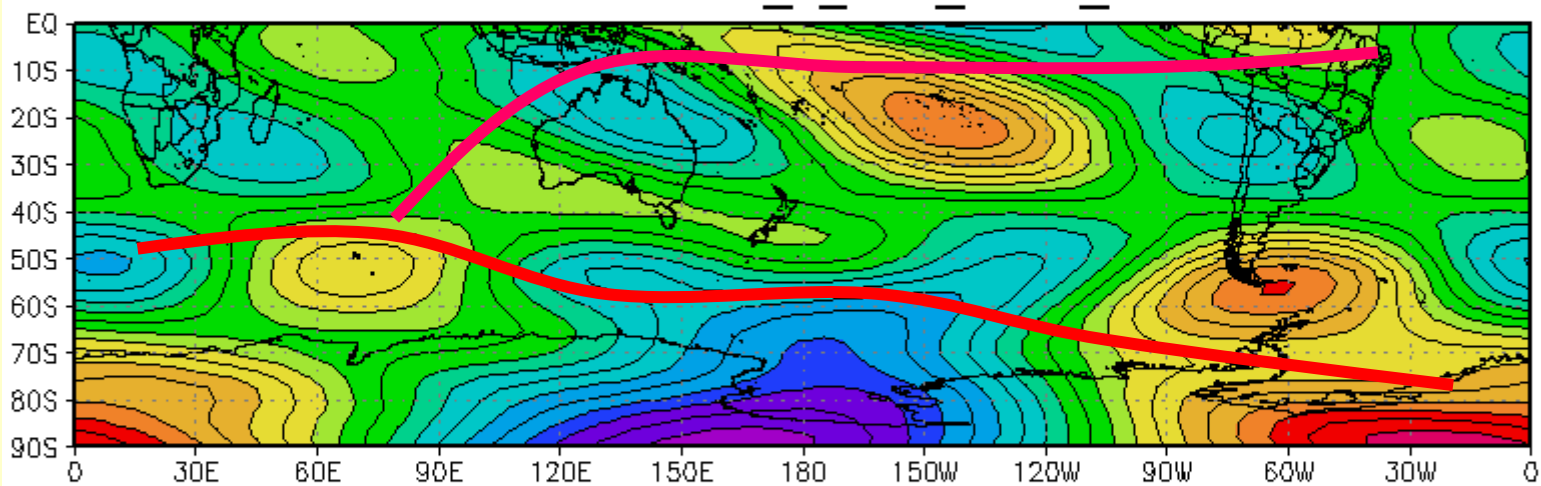
Zonal wind 200 mb - nw 5



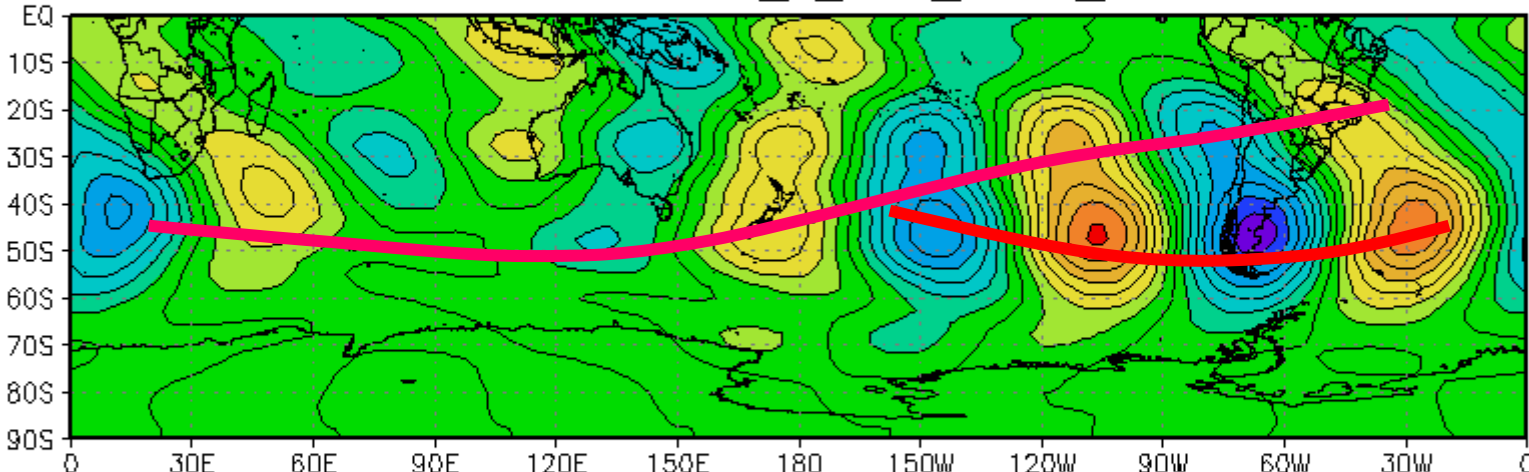
### Zonal Wind - M8\_U\_200\_0381\_WN=46



### Meridional Wind - M9\_V\_200\_0381\_WN=13



### Meridional Wind - M9\_V\_200\_0381\_WN=46



# **EUROBRISA... “downscaling” with RegCM3**

**COMPARISON BETWEEN CPTEC AND  
RegCM3 SEASONAL FORECASTS AND  
CLIMATOLOGY**

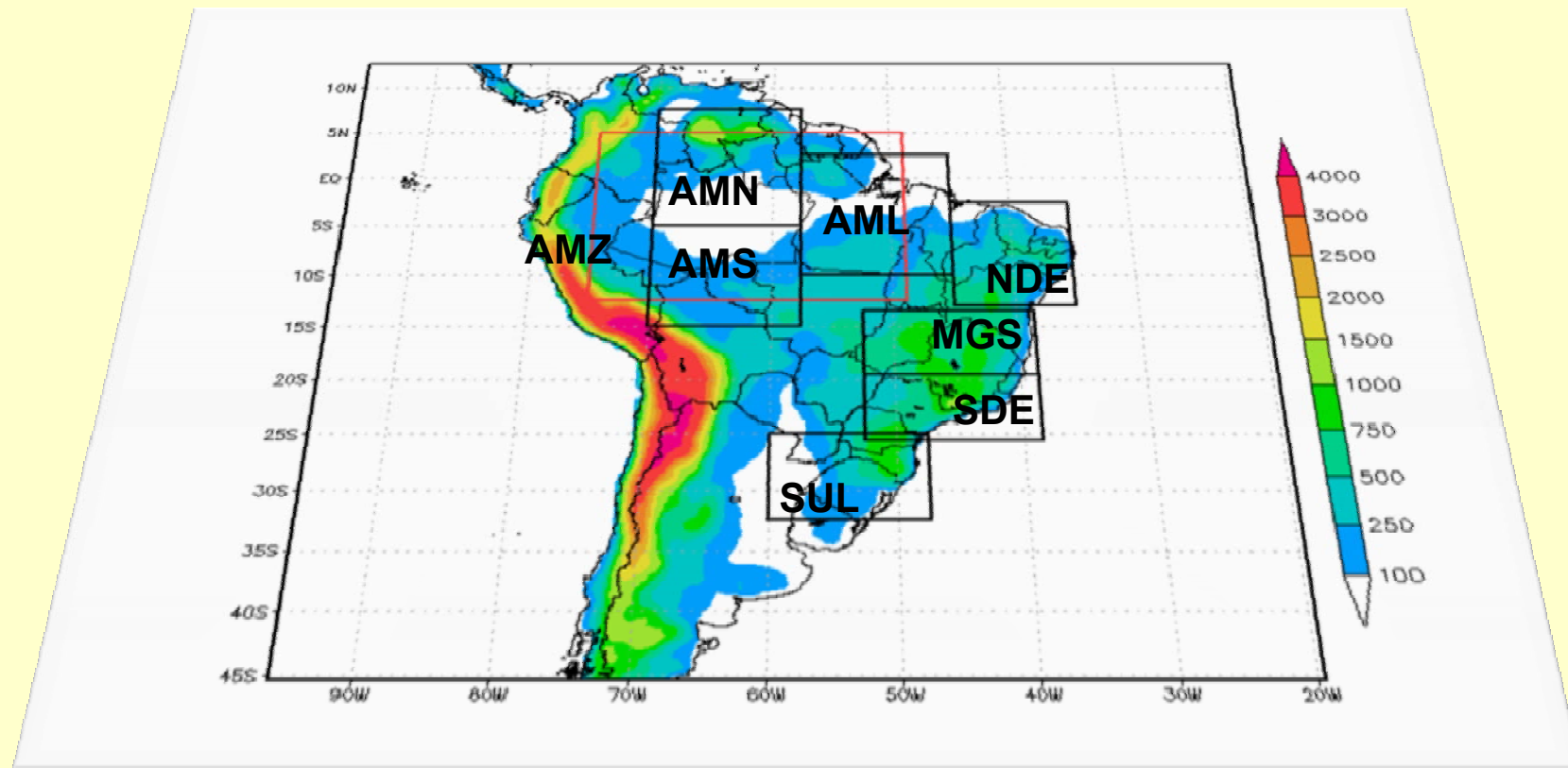
**RegCM3 was initialized with CPTEC  
SEASONAL MODEL from JJA/2005 up to  
ASO/2007**

# RegCM3 MODEL CONFIGURATION

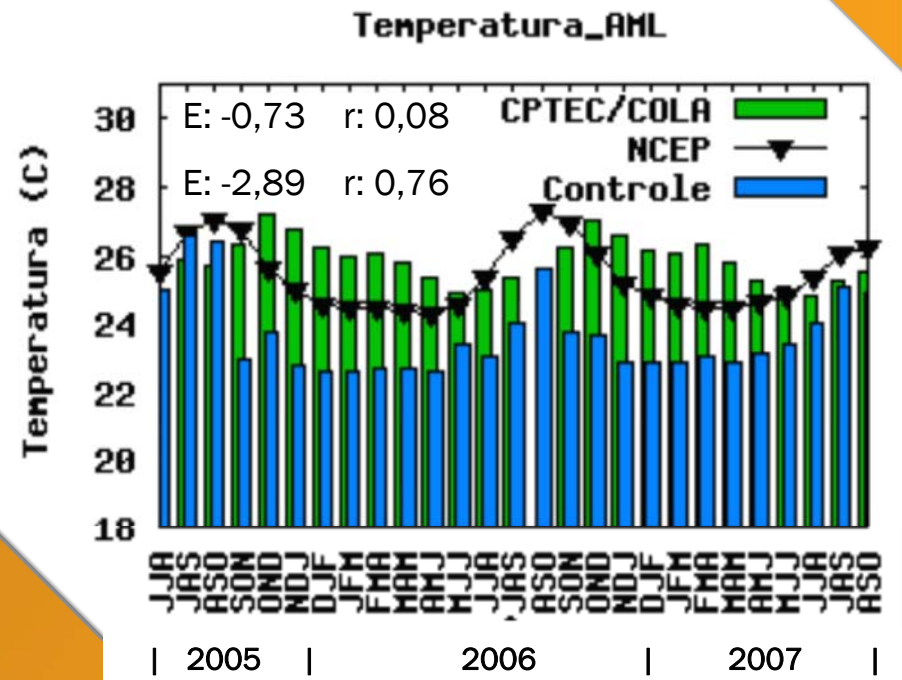
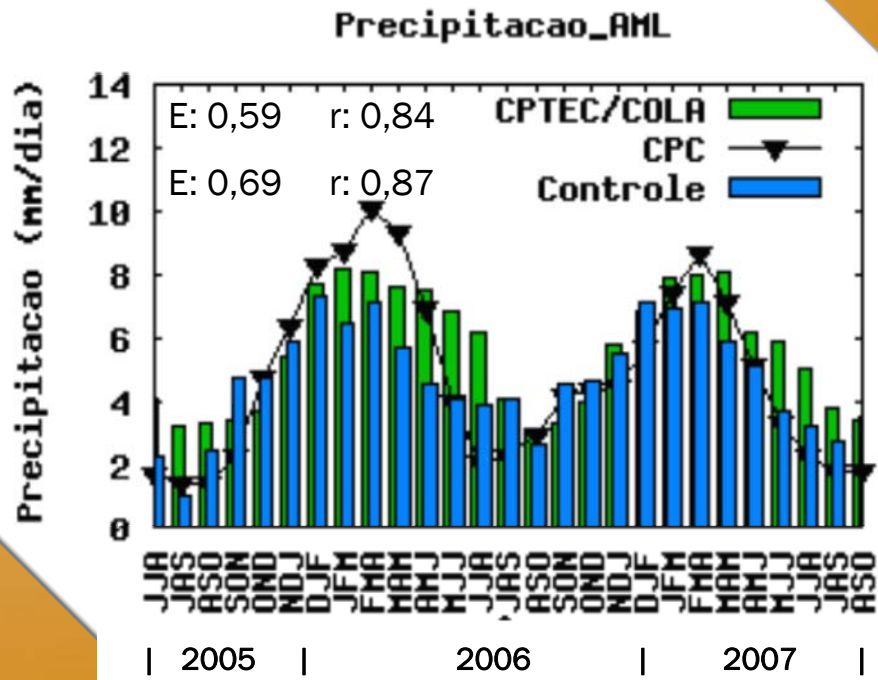
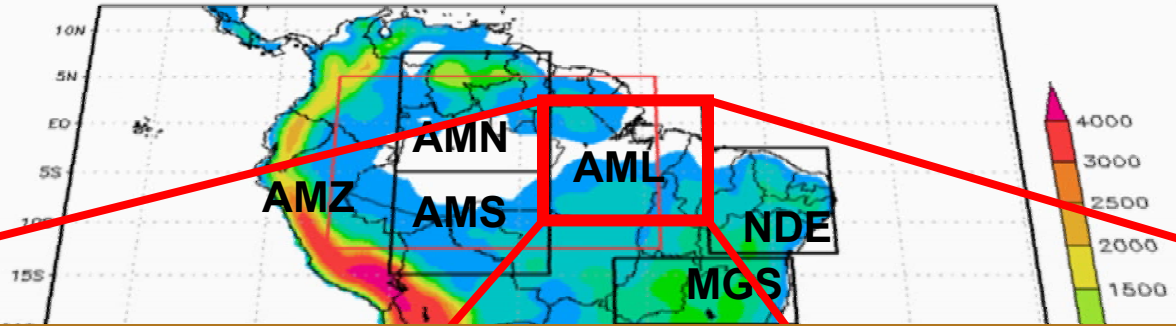
Characteristics		Configuration
Physics	Cumulus	Grell
	Radiation	CCM3
	Planetary Boundary Layer	Holstag
	Earth-Surface-Atmosphere	BATS
Vertical resolution		18 sigma levels
Horizontal resolution		60 km
Dynamics		Hidrostatic
Initial conditions and boundaries		CPTEC/COLA

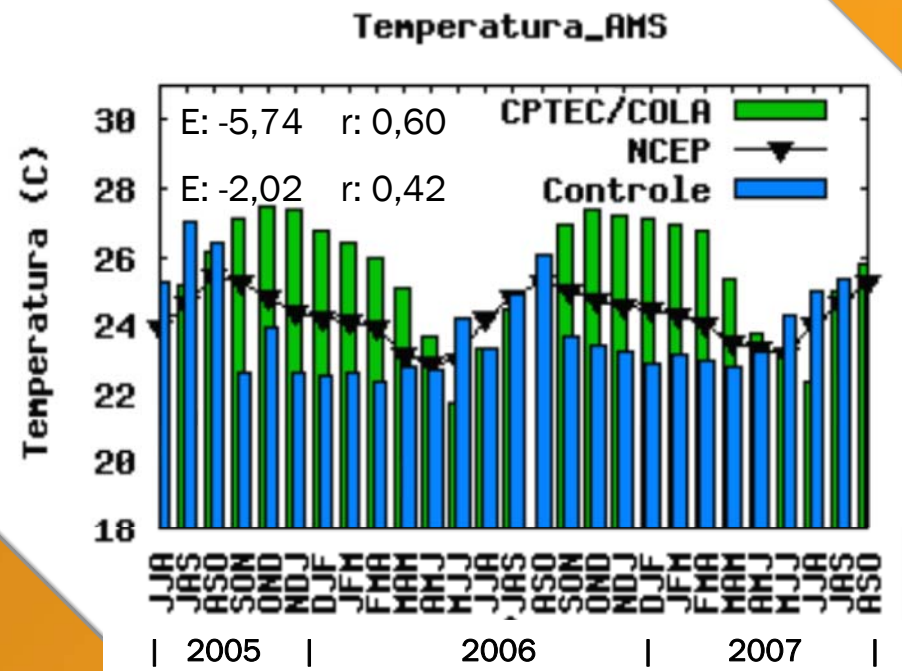
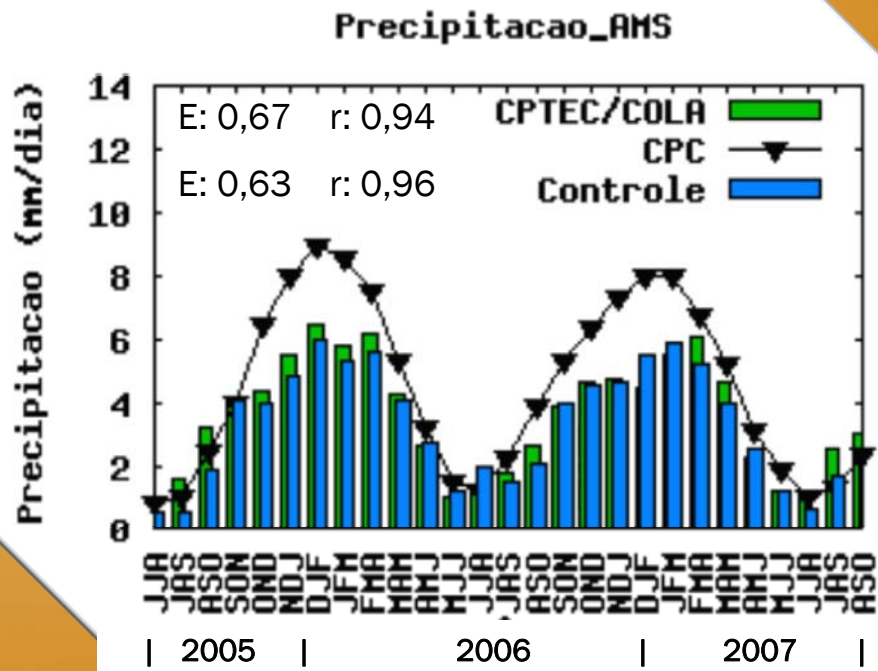
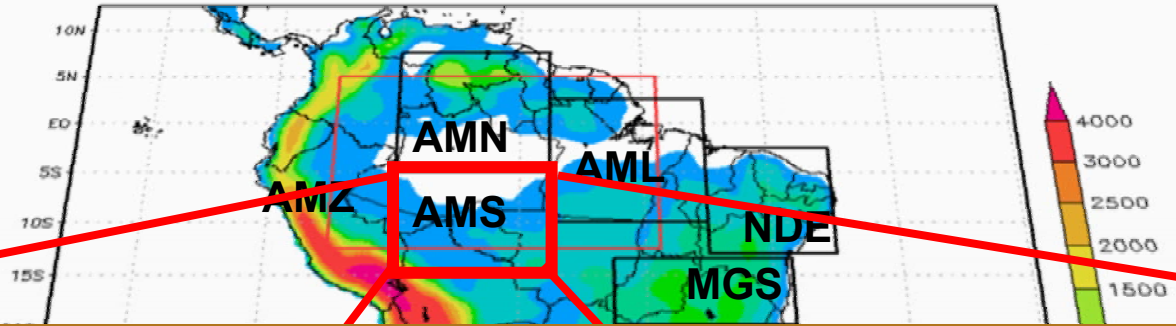
(Machado 2008)

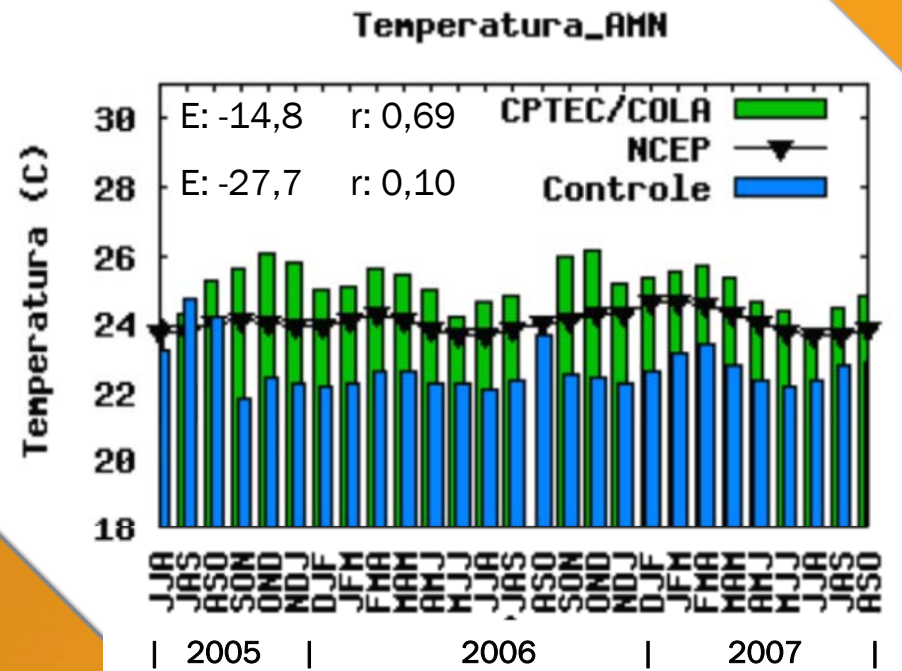
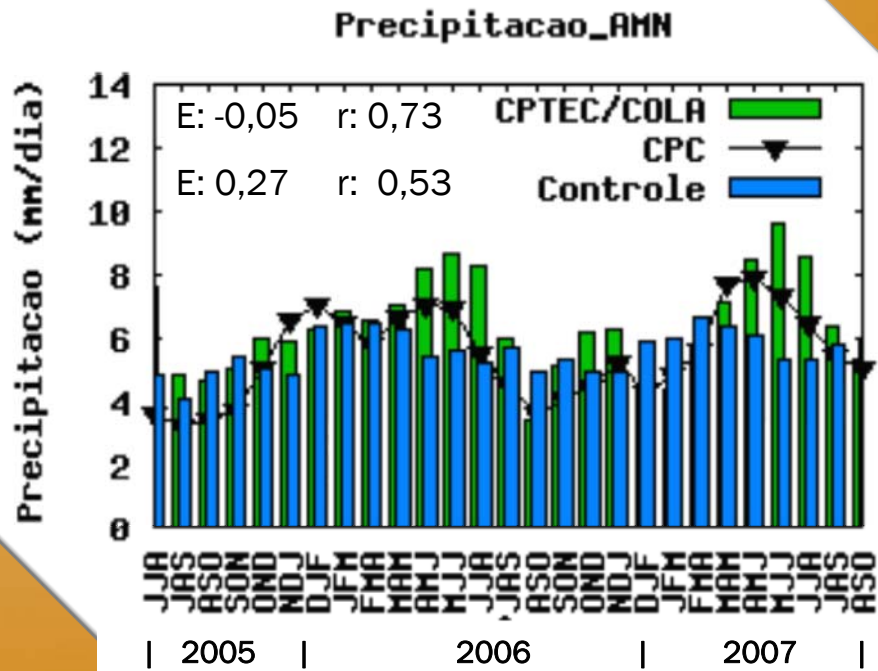
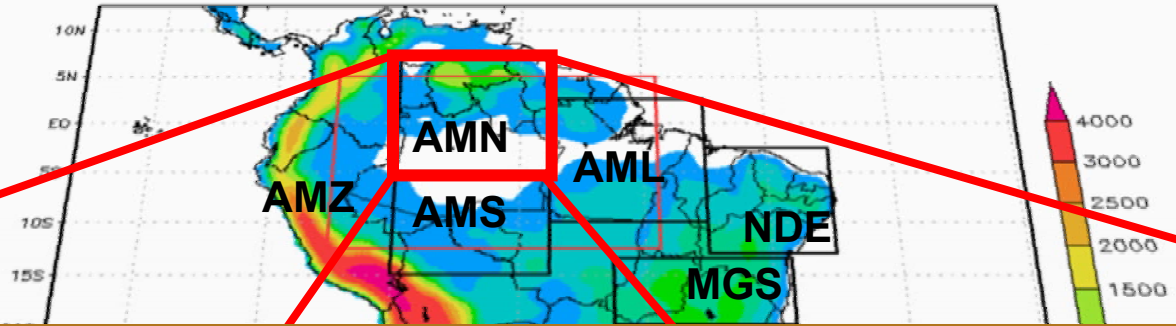
**Topography, Domain and 8 sub-regions where the 3 months forecast from RegCM3 were compared with the CPTC seasonal forecast model**

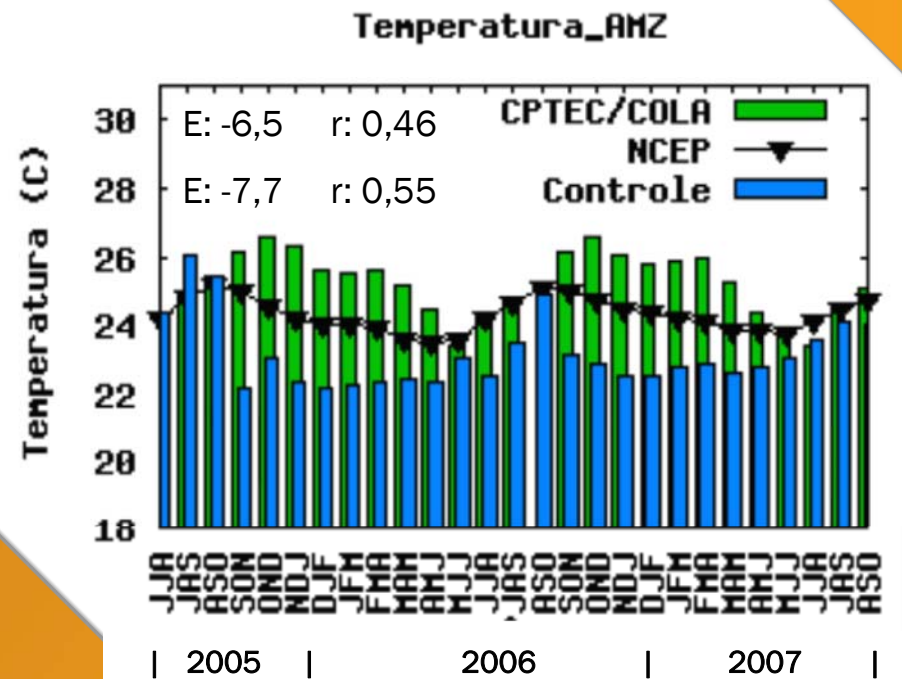
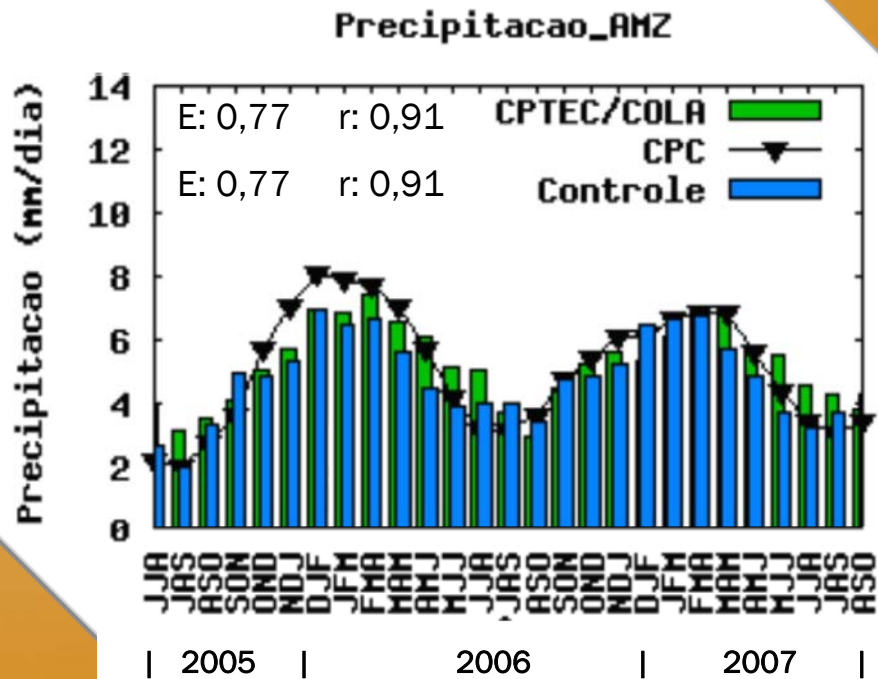
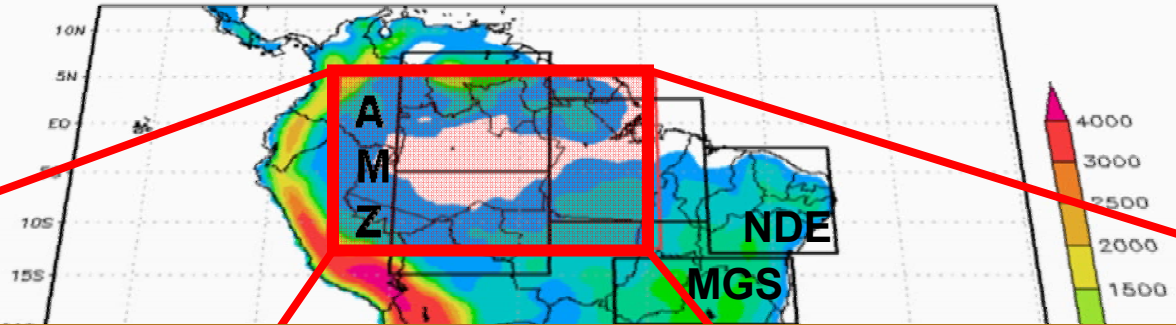


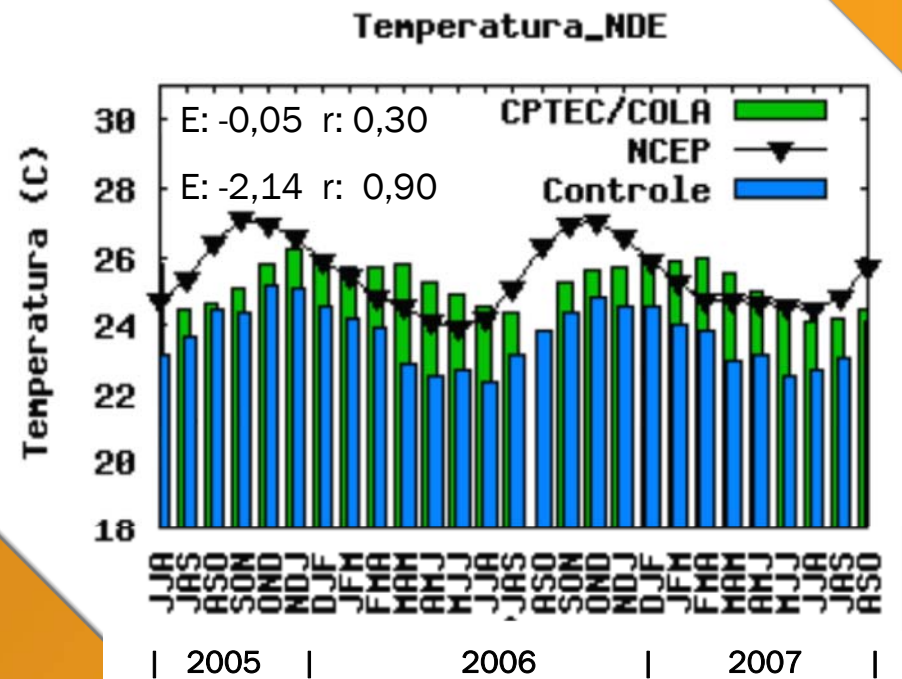
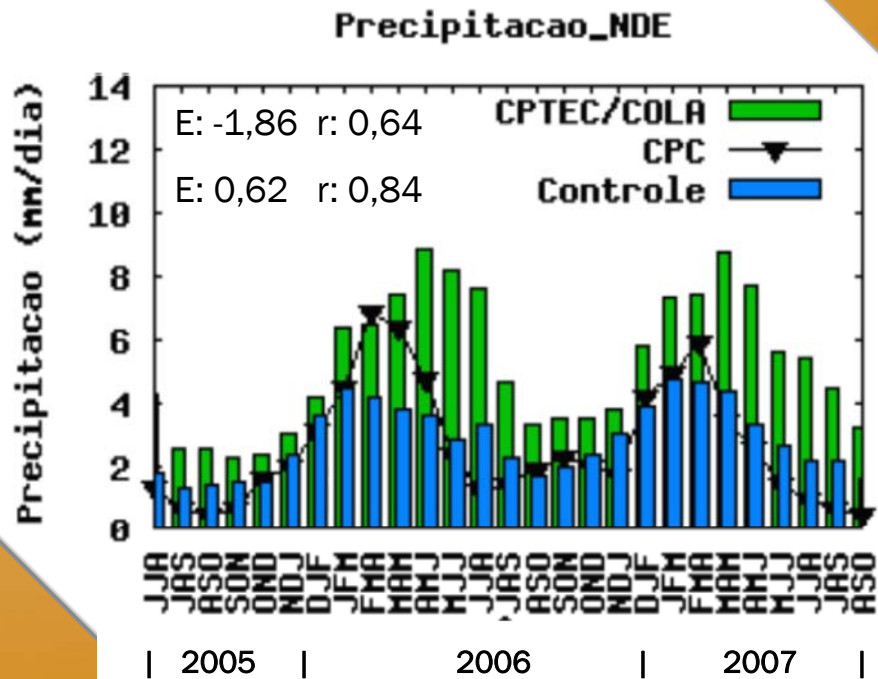
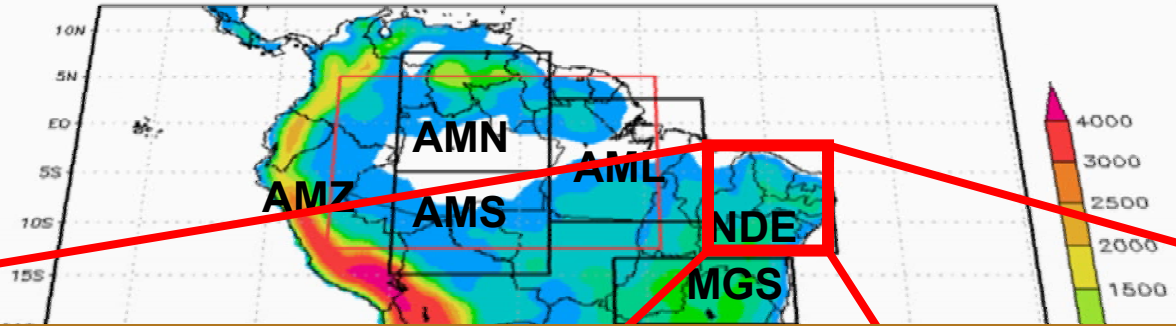
(Machado 2008)

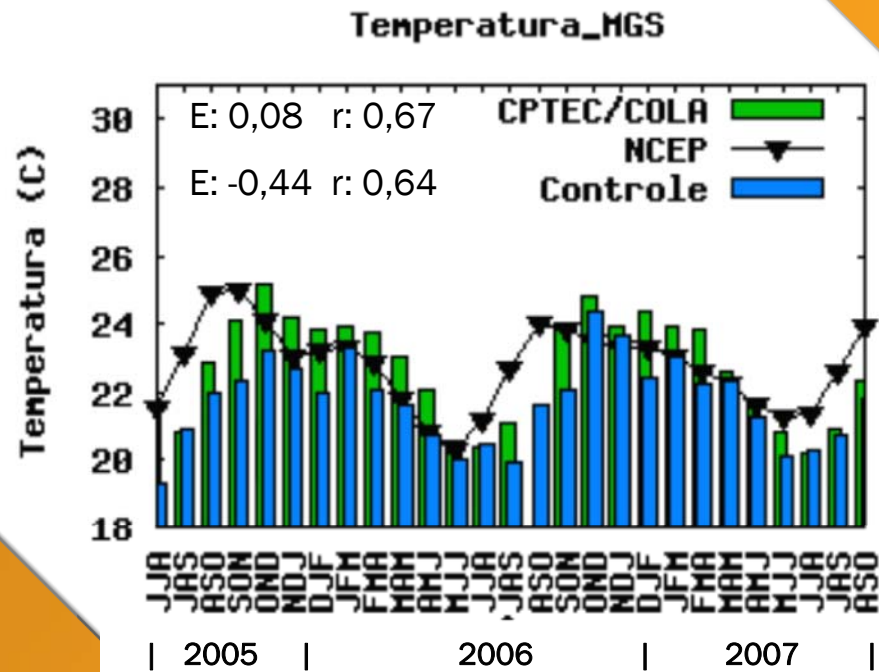
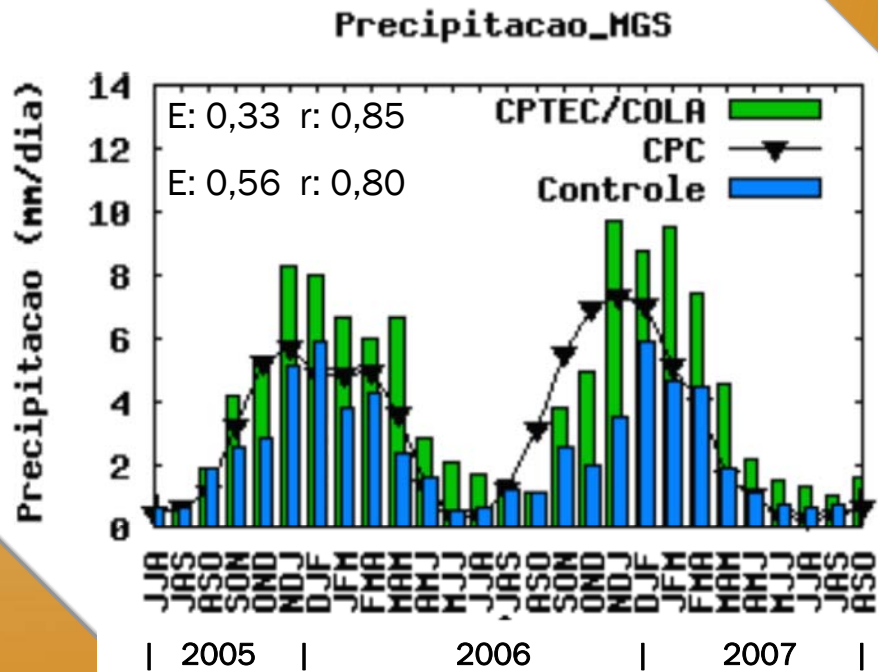
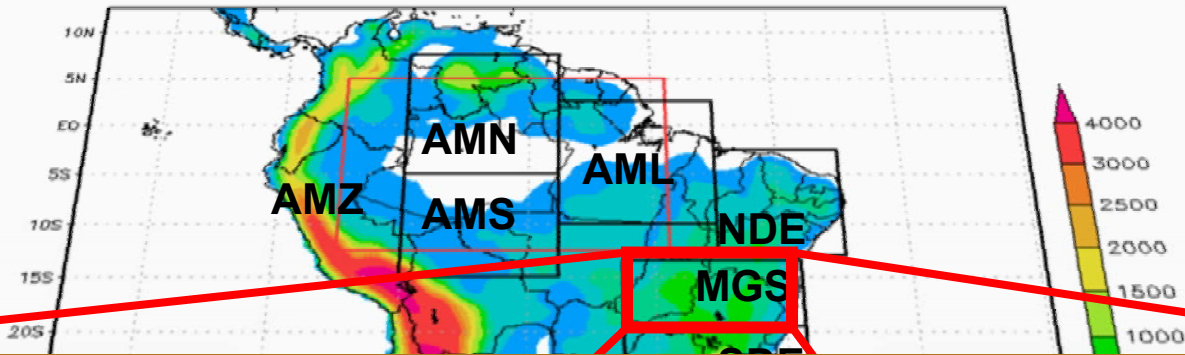


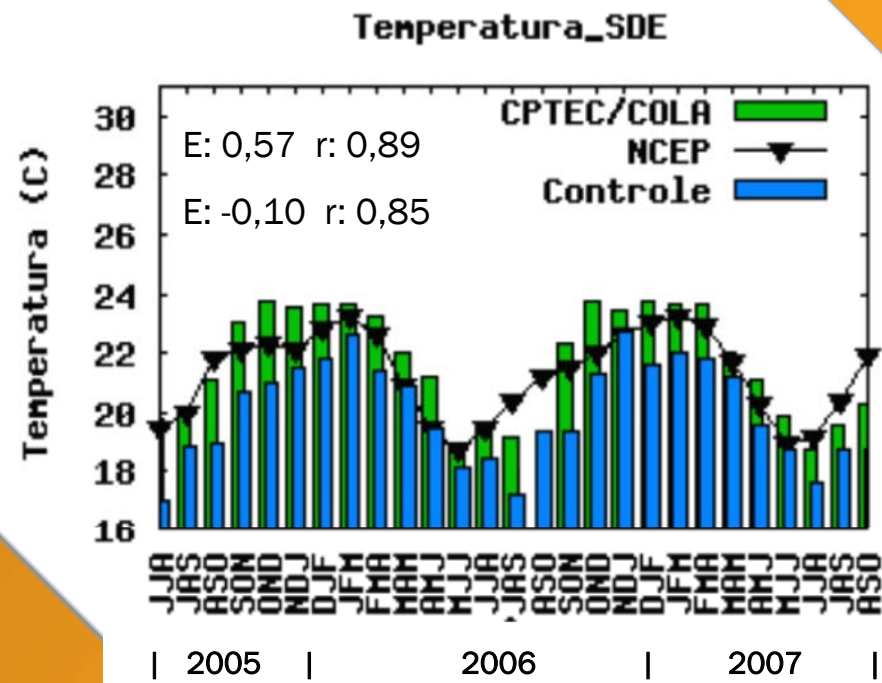
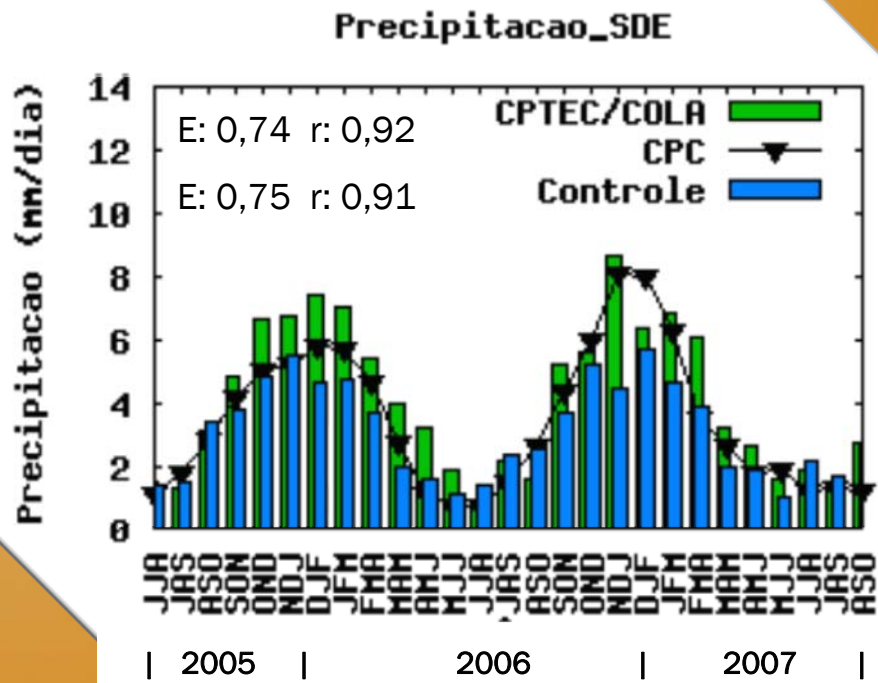
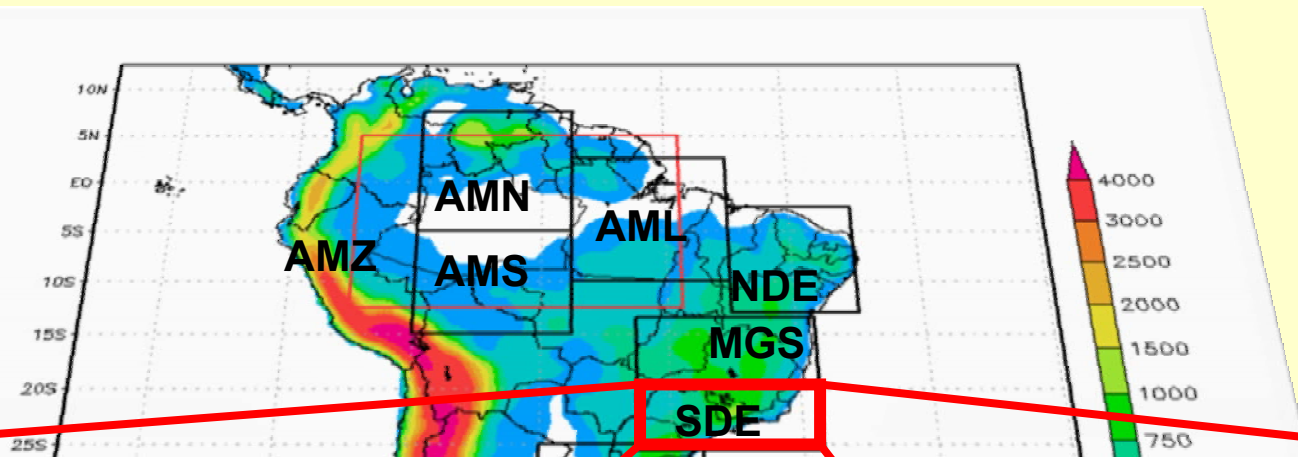


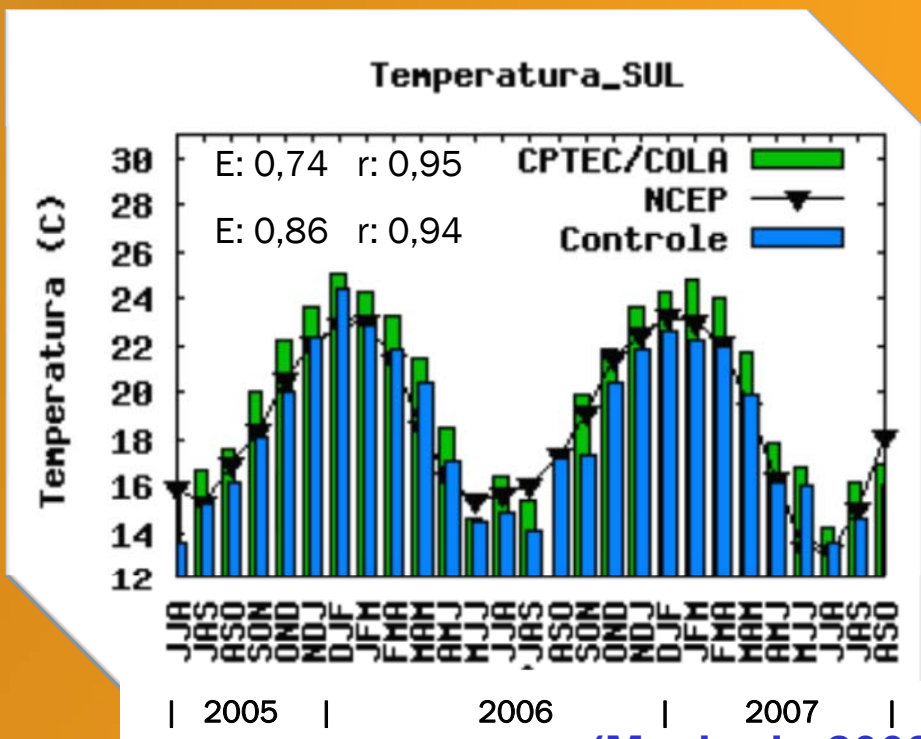
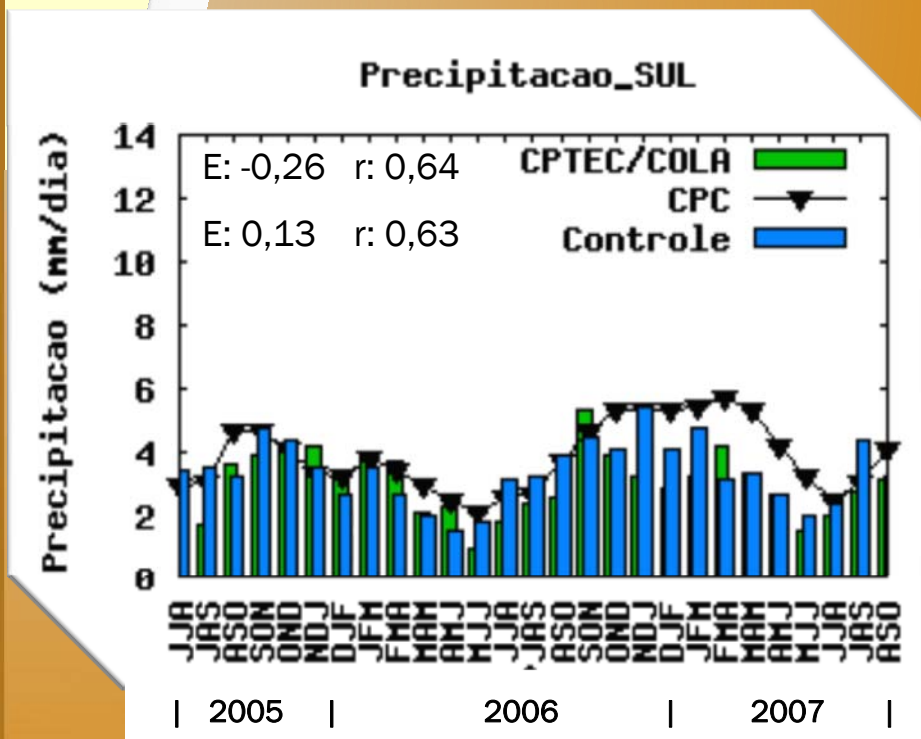
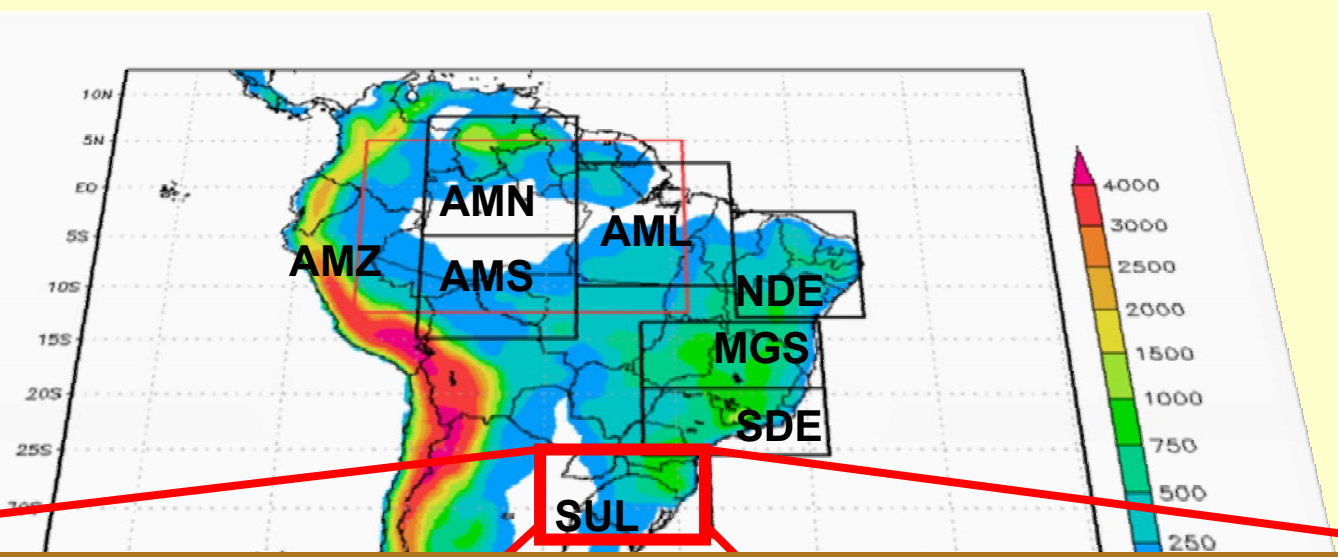












(Machado 2008)



**APEC Climate Center**

**Thanks to Dr. Saji N. Hameed and the APEC Climate  
Center for the invitation.**



GRUPO DE ESTUDOS CLIMÁTICOS

CLIMATE STUDIES GROUP

**THANK YOU FOR YOUR ATTENTION**

**GRACIAS POR SU ATENCIÓN**